



What can we learn from measurements of $^{13}\text{C}:^{12}\text{C}$ and $^{14}\text{C}:\text{C}$ in atmospheric CO_2 ?



... or, Other tracers that are hard to measure well and tell you about the carbon cycle

John B. Miller, NOAA Global Monitoring Laboratory

Outline

1. ^{14}C

1. Background
2. Signals in the data
3. Global inverse modeling
4. Urban

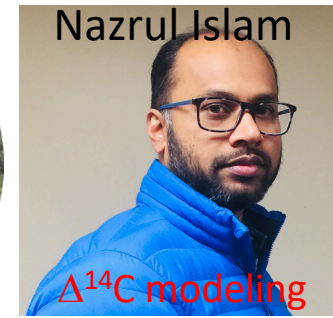
2. ^{13}C

1. Evolution of use
2. Region signals
3. Global signals

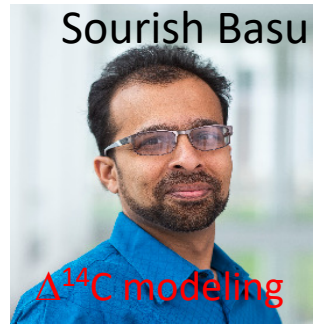
Scott Lehman



Nazrul Islam



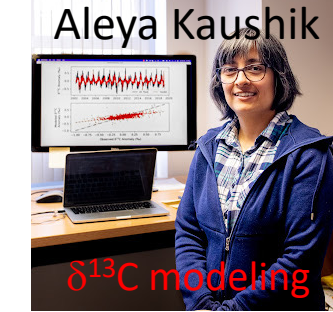
Sourish Basu



Sylvia Michel



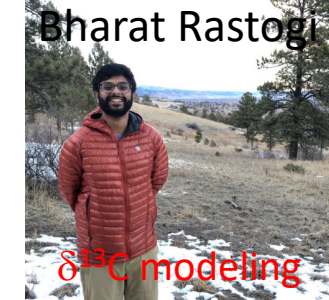
Aleya Kaushik



Jocelyn Turnbull



Bharat Rastogi



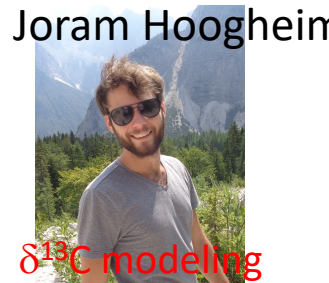
Ash Ballantyne



Caroline Alden



Joram Hoogheim



Wouter Peters



Ivar van der Velde



Radiocarbon basics

- ^{14}C is produced in the upper atmosphere and oxidized to CO_2 which is incorporated into the atmosphere-land-ocean system.
- Fossil fuels have no ^{14}C , because its half-life is ~ 6000 yrs. This makes $^{14}\text{CO}_2$ an ideal tracer for fossil CO_2 .
- As with stable isotopes, we report ^{14}C abundance ($\sim 10^{-12}$ of CO_2) using “delta” notation:

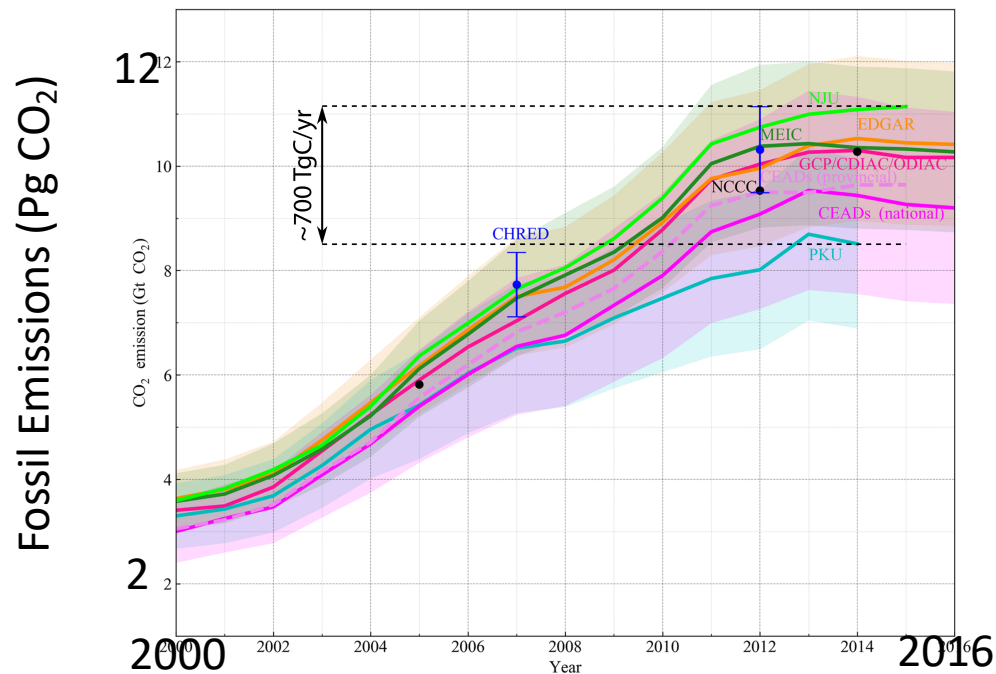
$$\Delta \approx \left(\frac{R_{\text{sample}}}{R_{\text{standard}}} - 1 \right) \times 1000 \text{ ‰}$$

→ as Delta decreases, $^{14}\text{C}:\text{C}$ decreases, indicating more fossil

Why Top-Down $^{14}\text{CO}_2$?

Significant spread in national inventories.

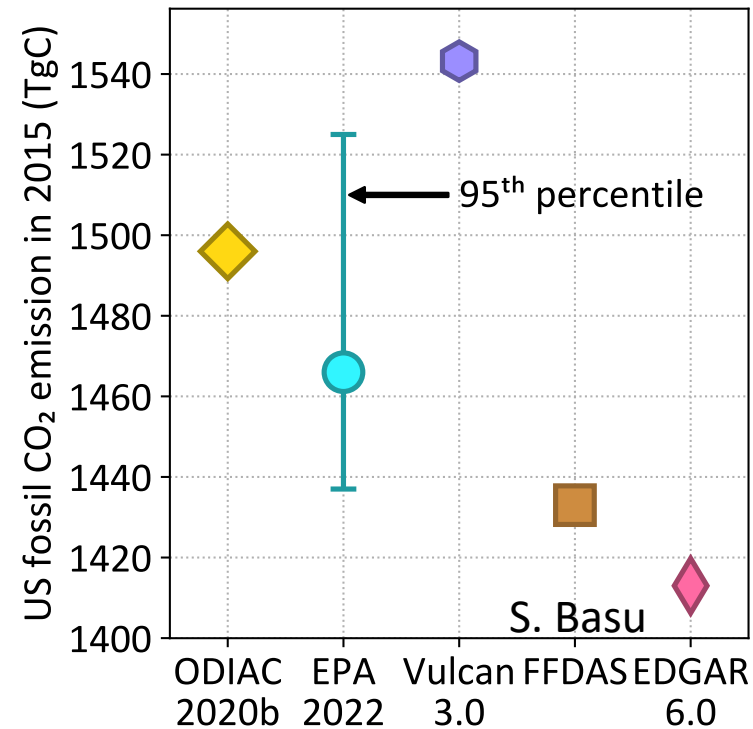
China: ~ 700 Tg C/yr



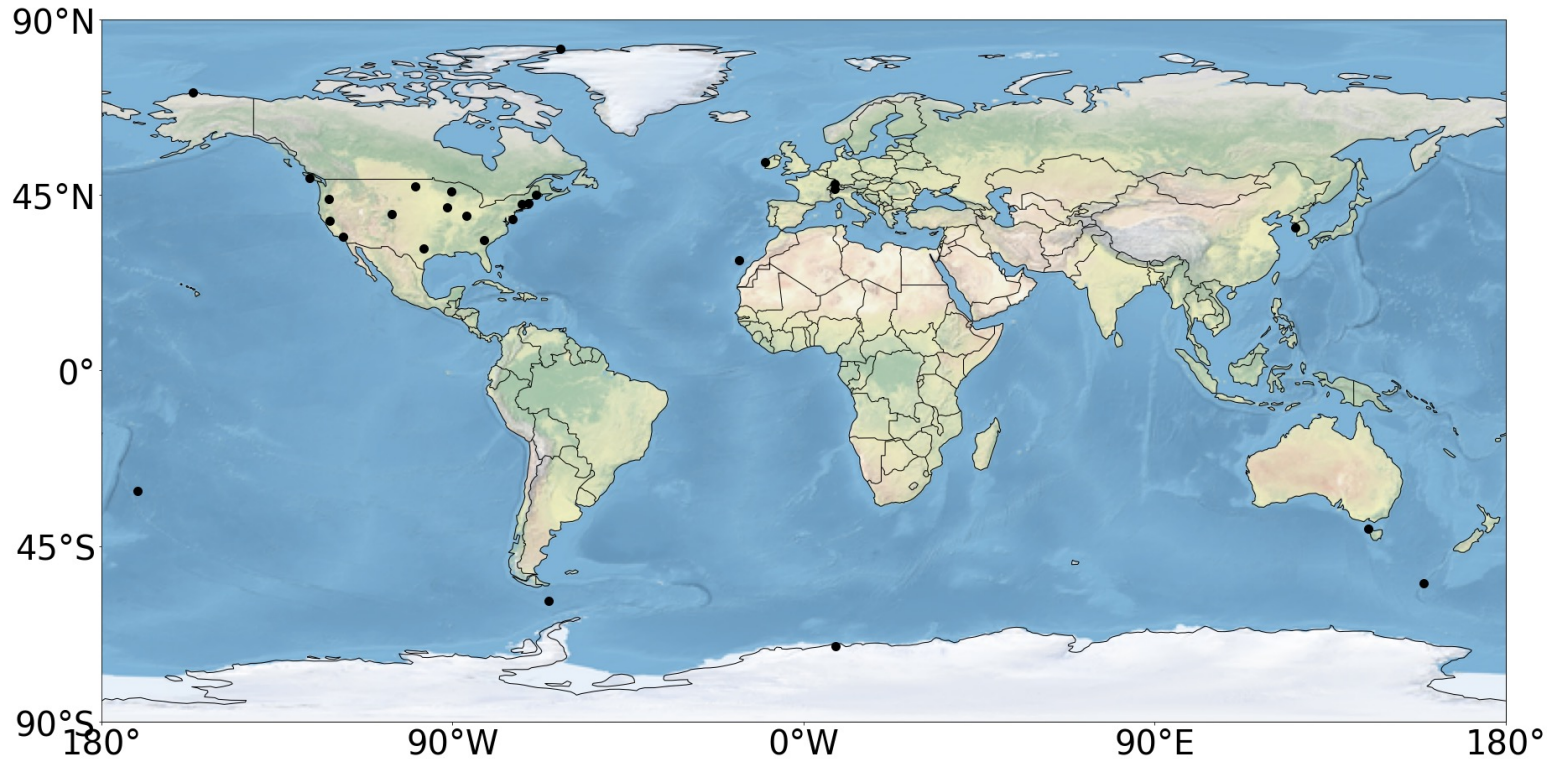
Evaluating China's fossil-fuel CO₂ emissions from a comprehensive dataset of nine inventories

Pengfei Han¹, Ning Zeng², Tom Oda³, Xiaohui Lin⁴, Monica Crippa⁵, Dabo Guan^{6,7}, Greet Janssens-Maenhout⁵, Xiaolin Ma⁸, Zhu Liu^{6,9}, Yuli Shan¹⁰, Shu Tao¹¹, Haikun Wang⁸, Rong Wang^{11,12}, Lin Wu⁴, Xiao Yun¹¹, Qiang Zhang¹³, Fang Zhao¹⁴, and Bo Zheng¹⁵

USA: ~130 Tg C/yr

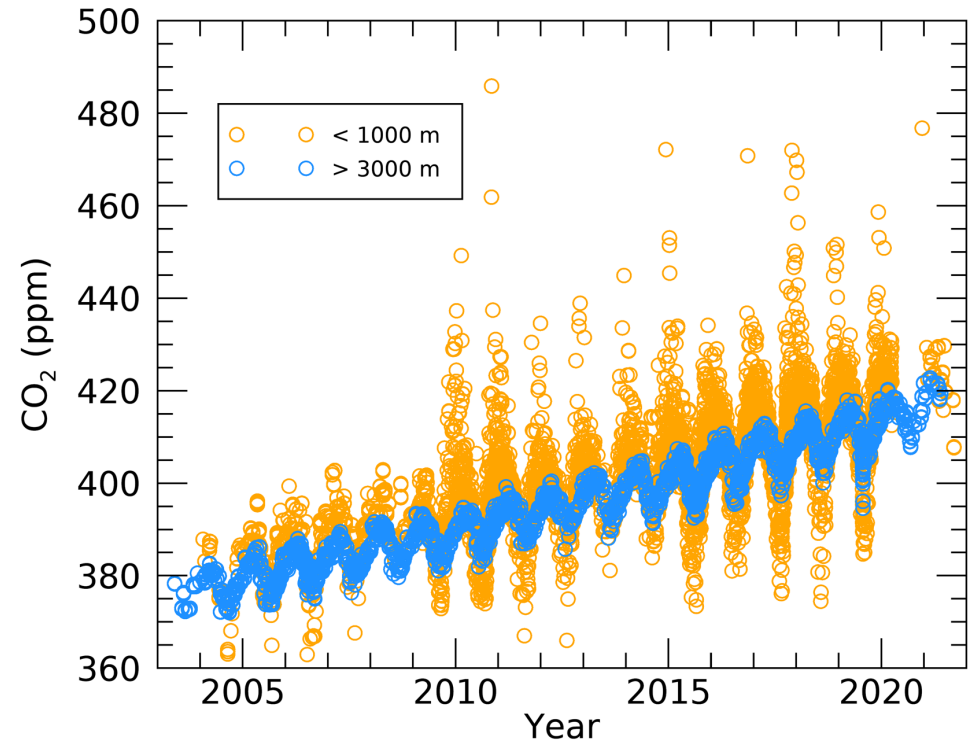
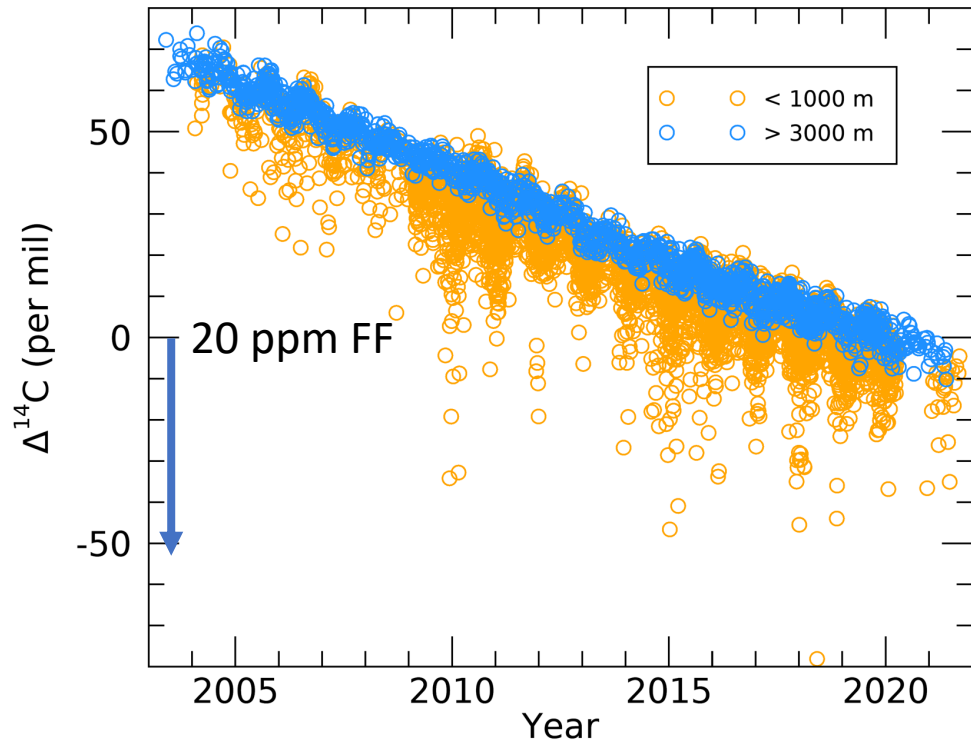


Measurement sites are global but mainly in North America

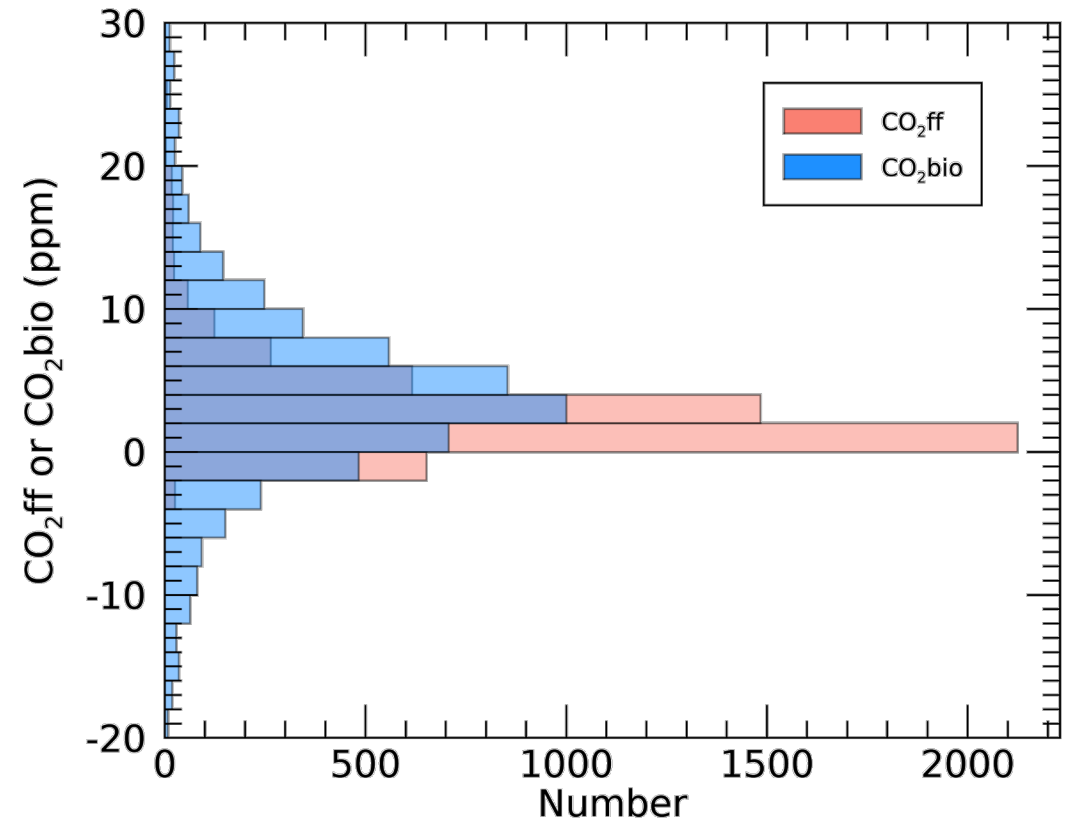
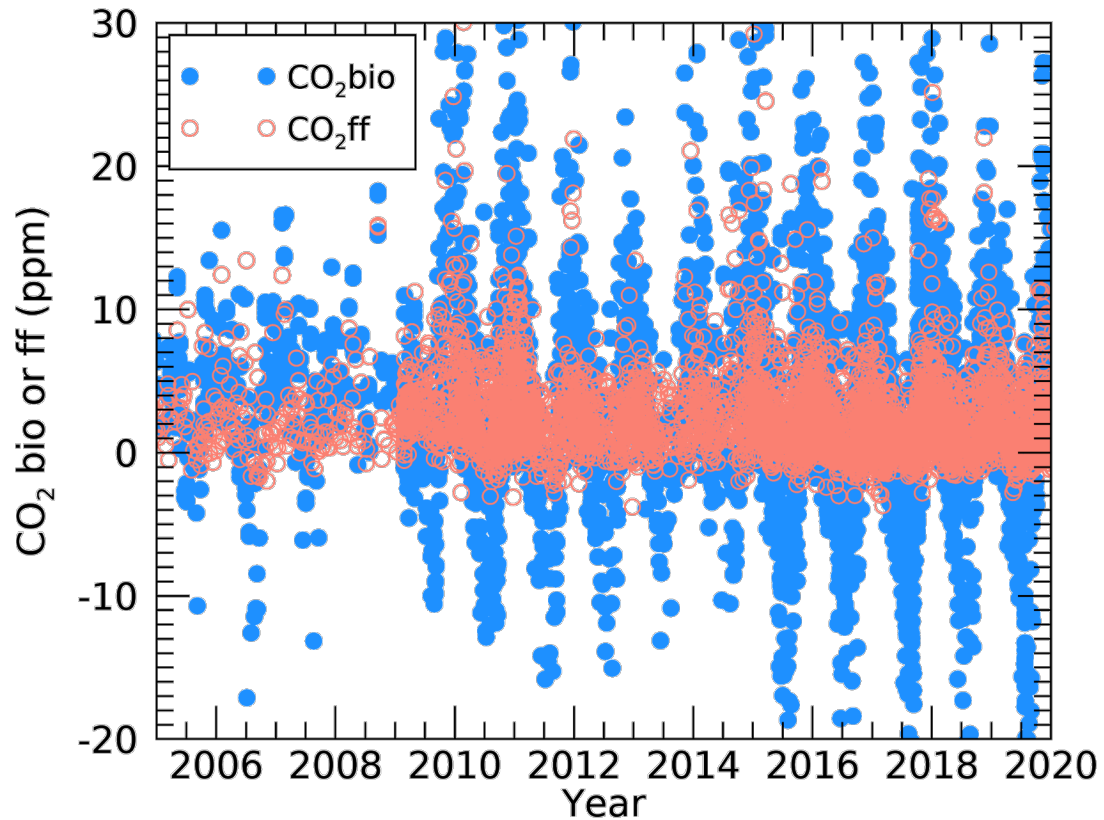


- ~ 1000 measurements per year.
- US sites 1 – 3 x per week.
- Most from U.S. tall towers (NOAA)
- Also aircraft sites
- Many global background sites from U. Heidelberg.

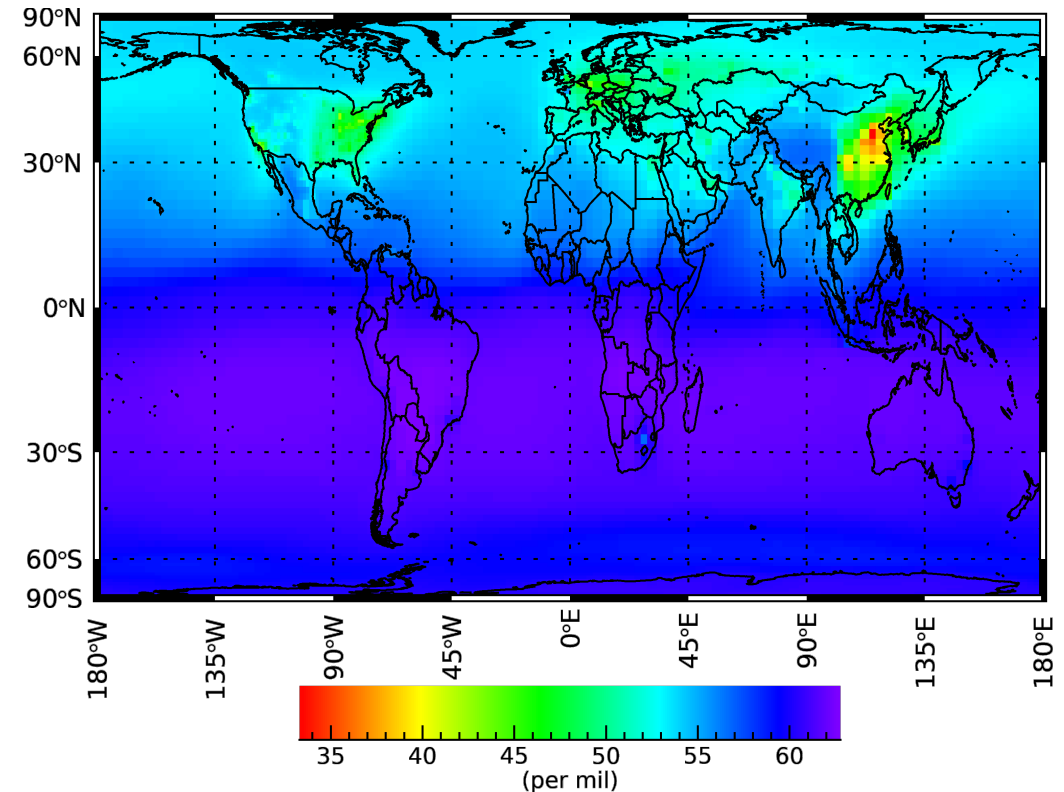
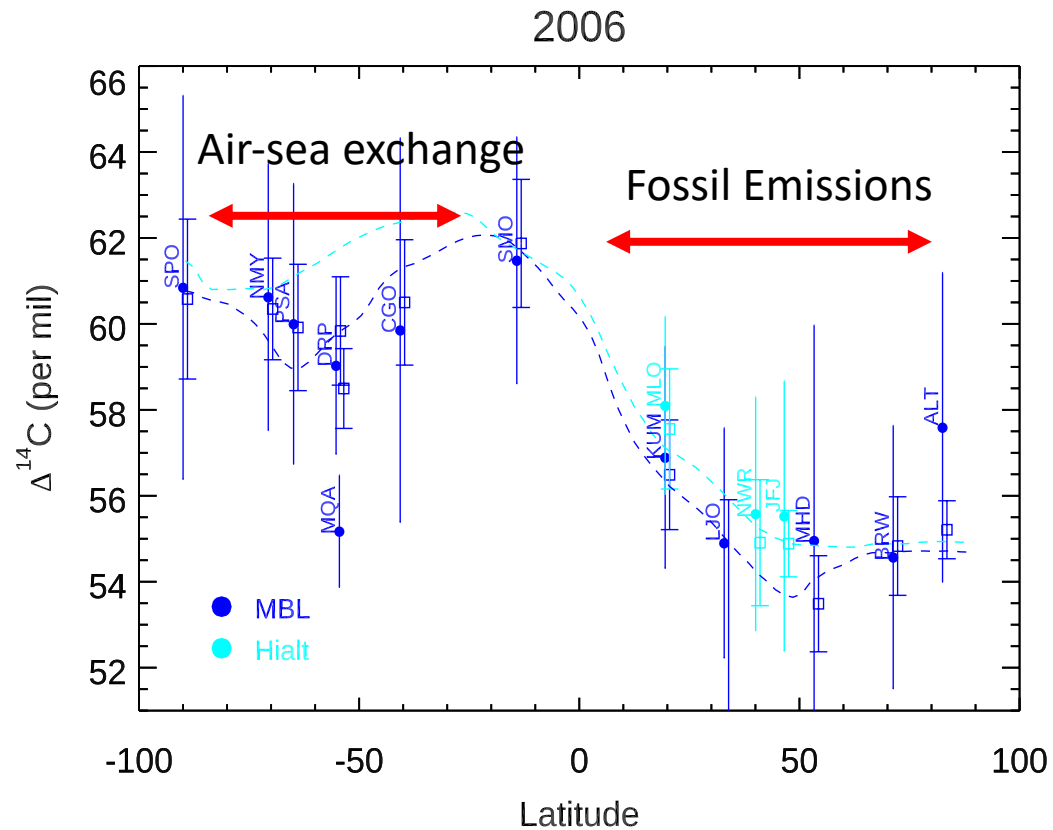
$\Delta^{14}\text{C}$ and CO_2 variations over U.S. are dominated by biology and fossil emissions, respectively



$\Delta^{14}\text{C}$ can separate fossil and biological components of continental CO_2

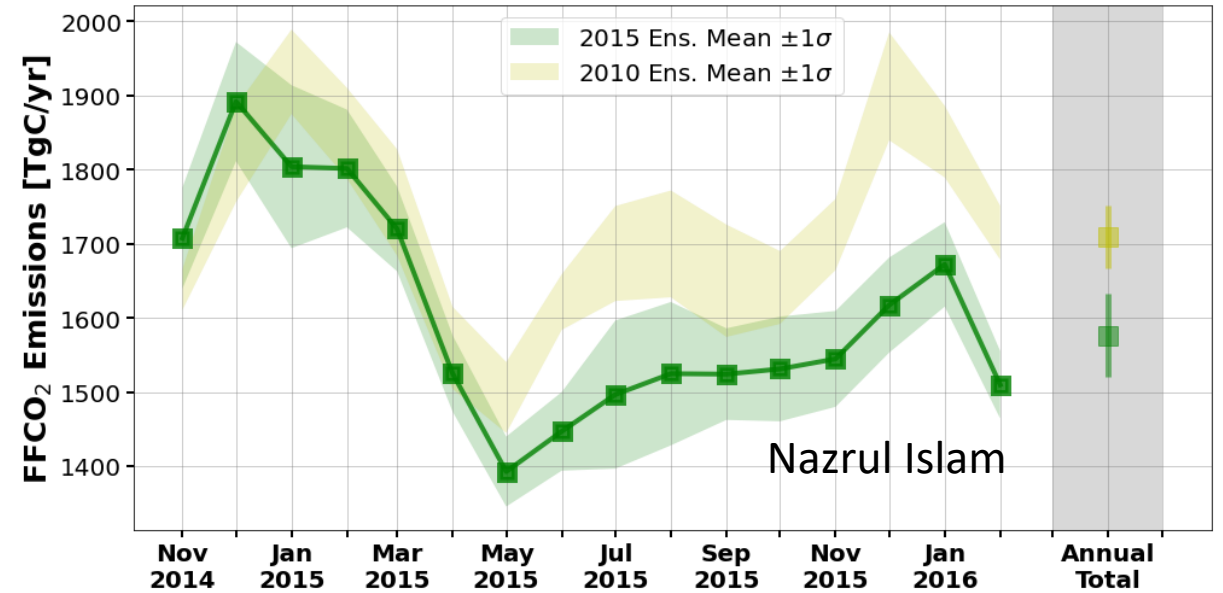
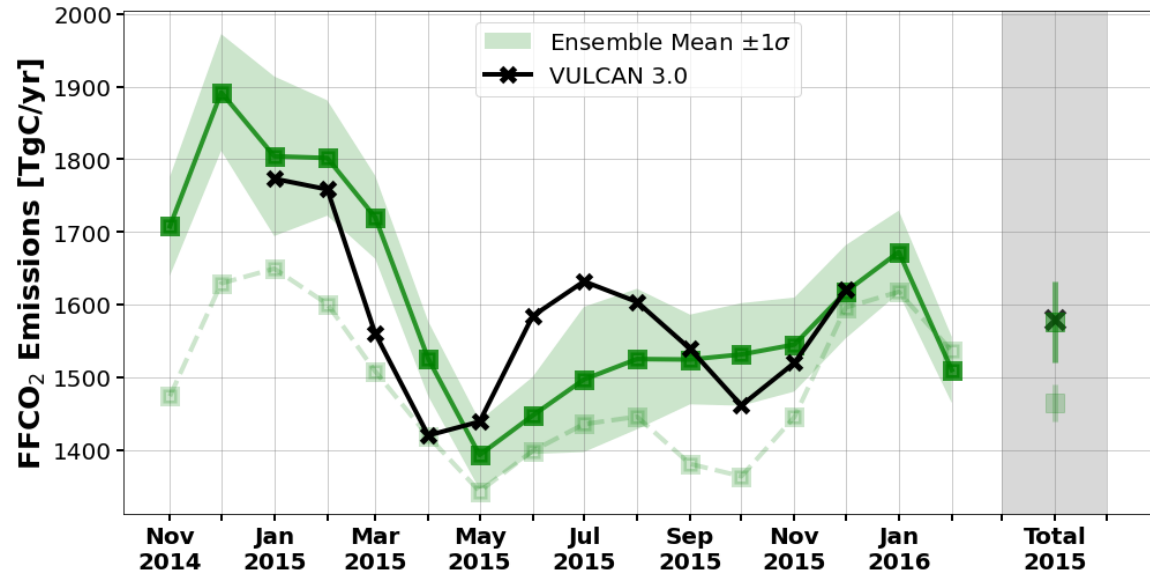


Forward model comparisons with data gives confidence ahead of inversion: 1) large scale lat. gradient



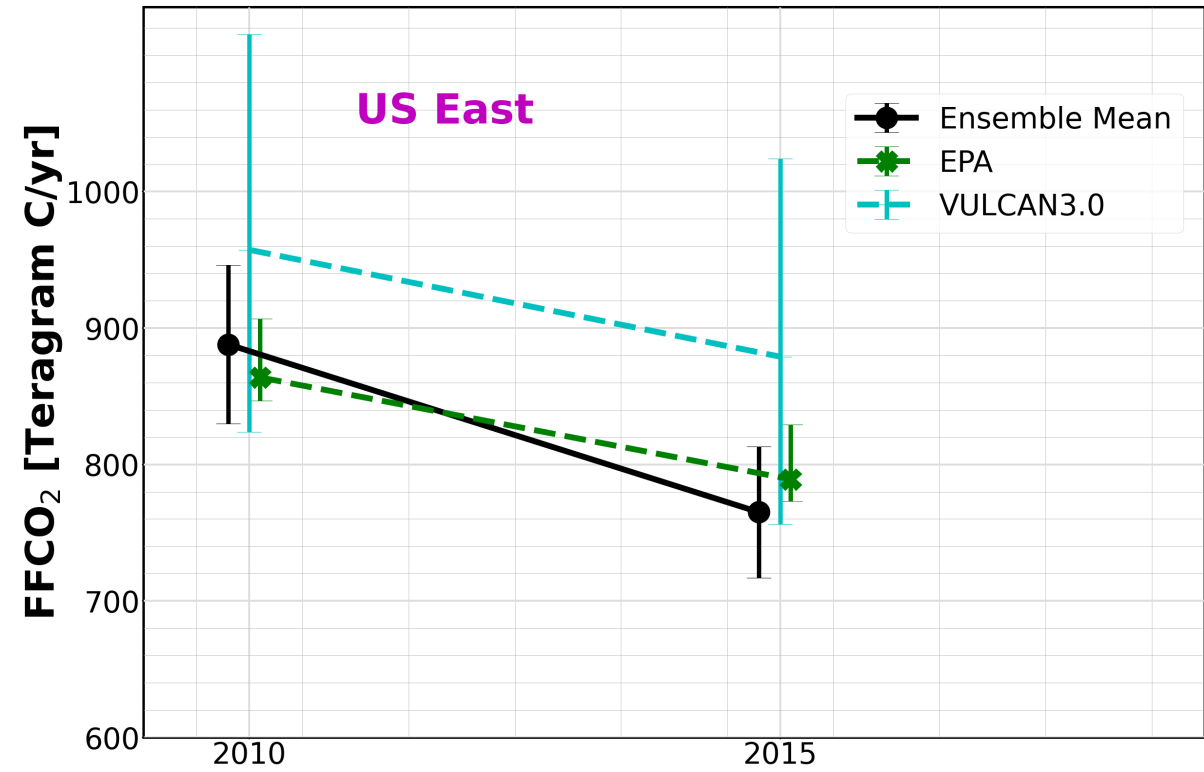
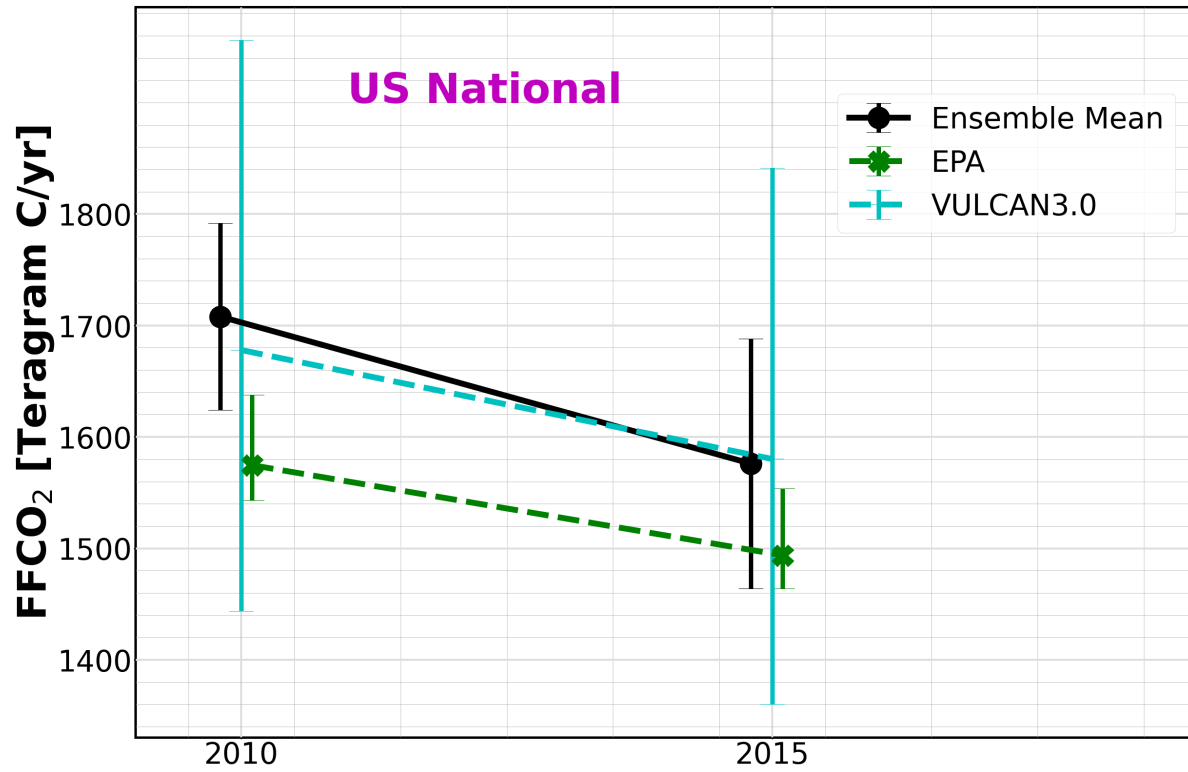
→ Model includes all $\Delta^{14}\text{C}$ budget terms: **Fossil**, cosmogenic and nuclear reactor production, **land and ocean disequilibrium**

US top-down fossil emissions for 2010 and 2015



Nazrul Islam

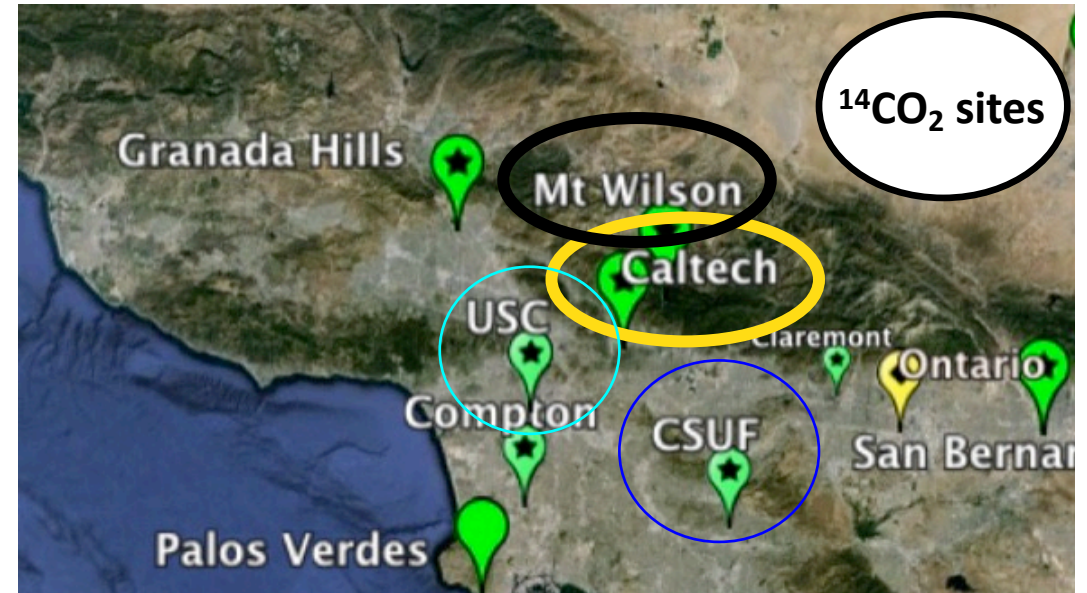
Top-down estimates show reductions consistent with US EPA (bottom-up)



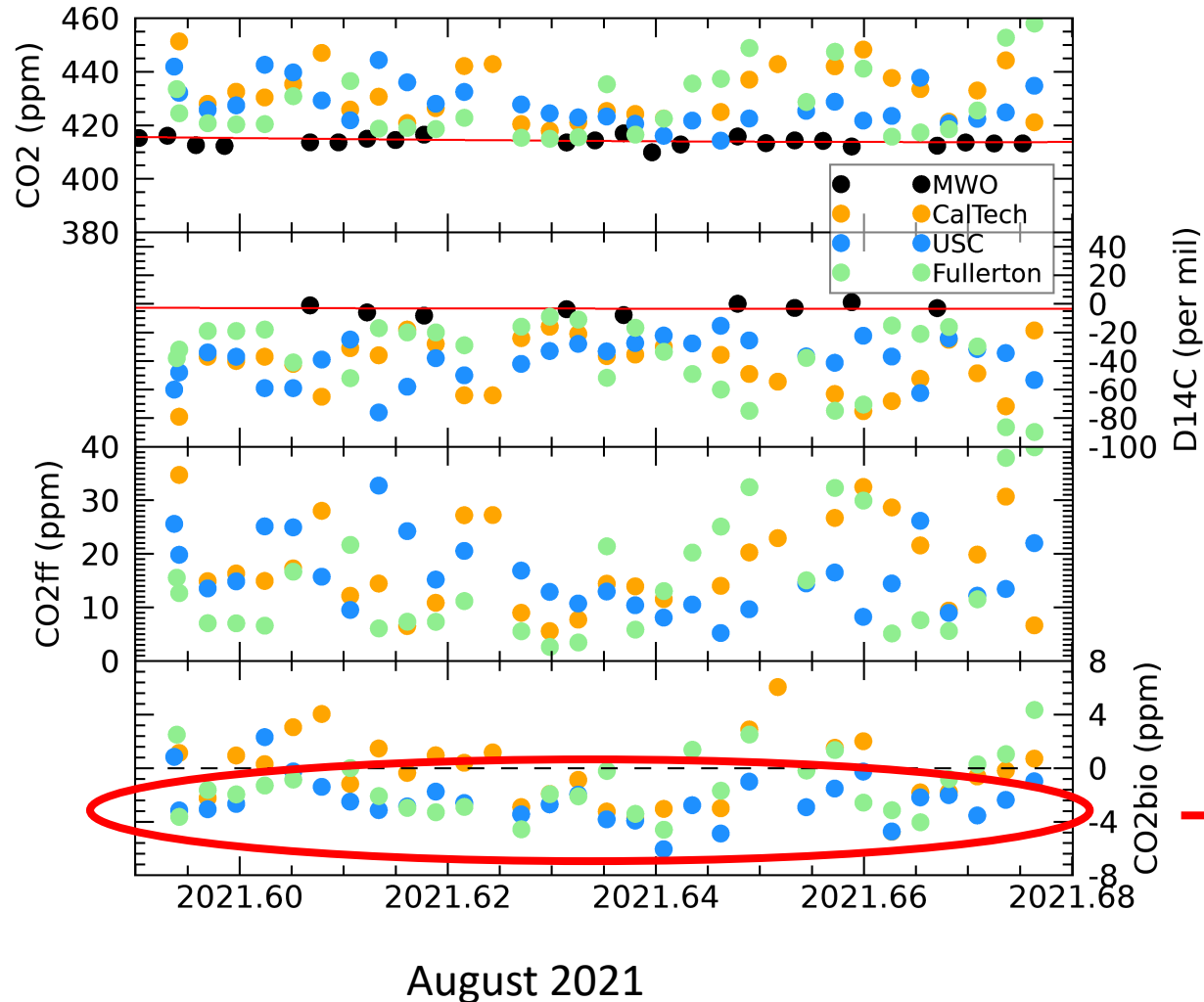
LA Megacity ^{14}C sampling

1. Quantify fossil (C_{ff}) and biospheric (C_{bio}) contributions to CO_2 enhancements (C_{xs})
2. 2015 sampling showed that C_{bio} seasonality was correlated with urban water use (PNAS, 2020).
3. Use C_{ff} to evaluate urban bottom-up fossil flux estimates.

LA Megacity GHG Network



Significant C_{bio} during SUNVEX-LA 2021



$$C_{obs} = C_{bg} + C_{bio} + C_{ff}$$

CalTech (pointing to C_{obs})

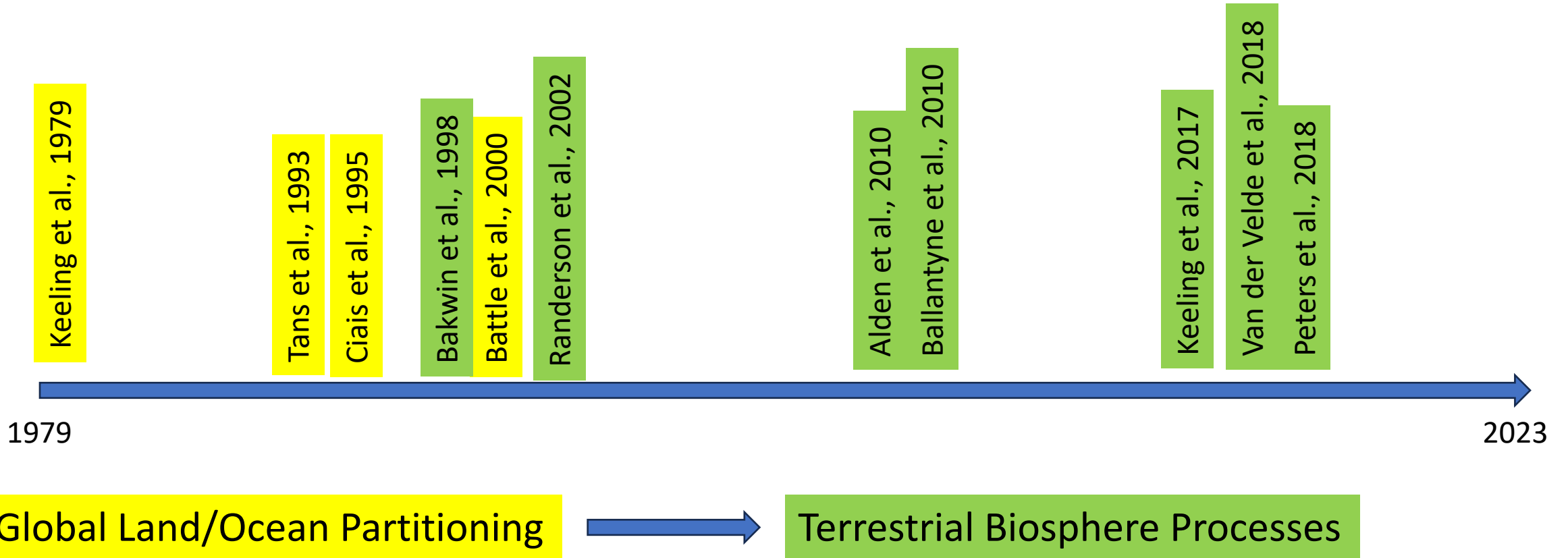
^{14}C (pointing to C_{ff})

MWO (pointing to C_{bio})

$$C_{xs} = C_{bio} + C_{ff}$$

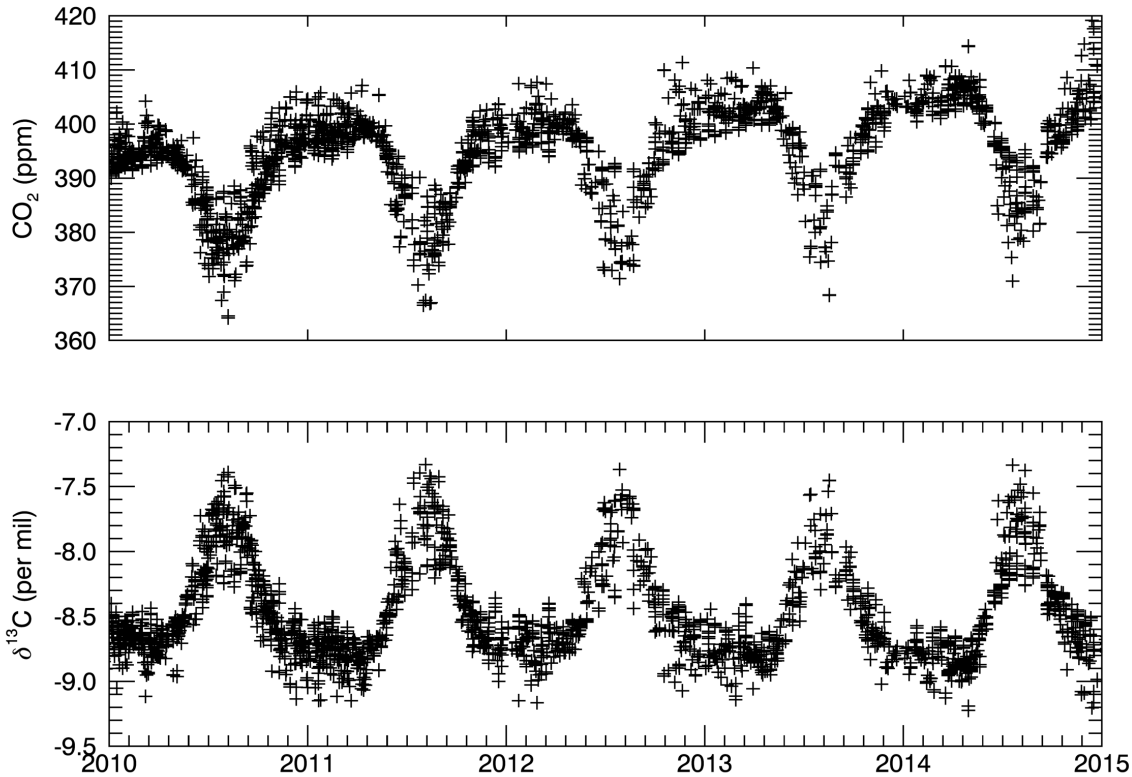
August CO₂ uptake consistent with 2015 results published in Miller et al., 2020 PNAS.

Atmospheric $^{13}\text{C}\text{CO}_2$ in transition...

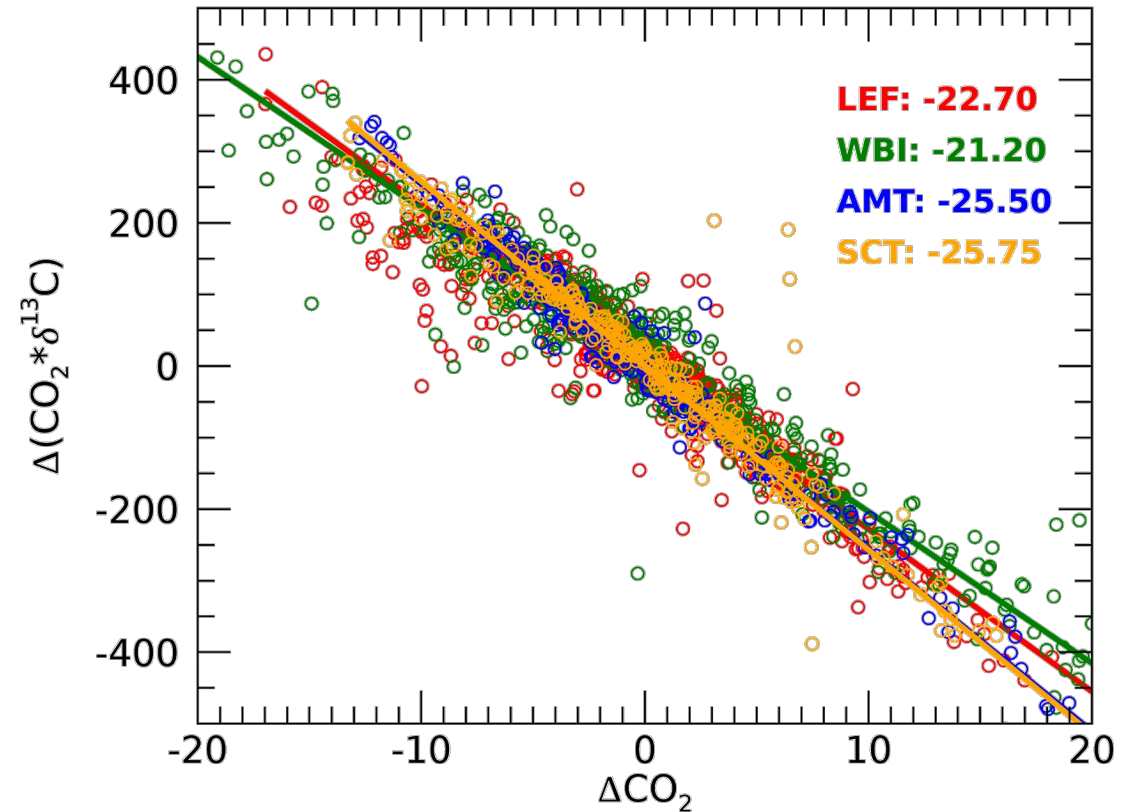


CO₂ and ¹³C:¹²C are strongly connected via plant photosynthesis.

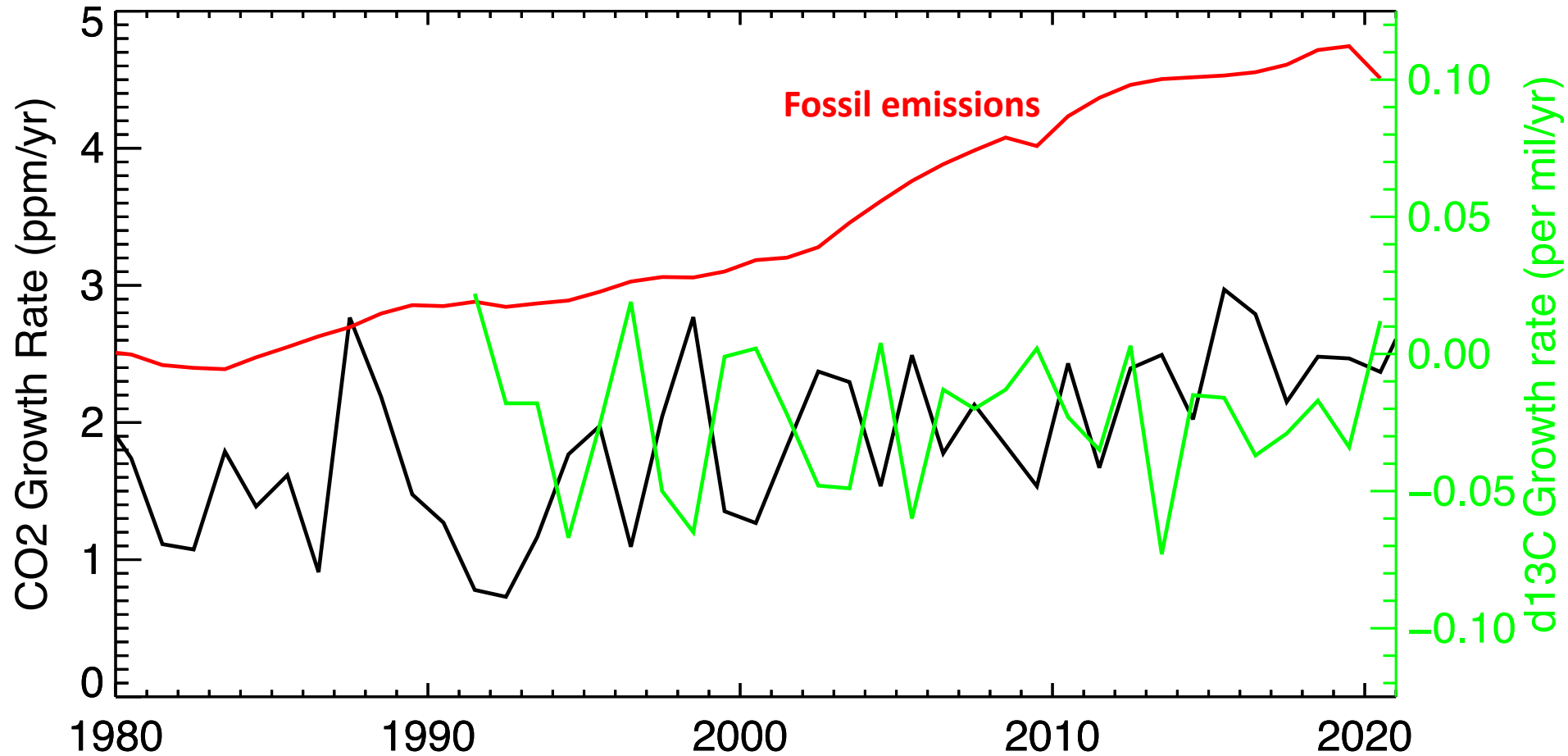
LEF tall tower, northern Wisconsin



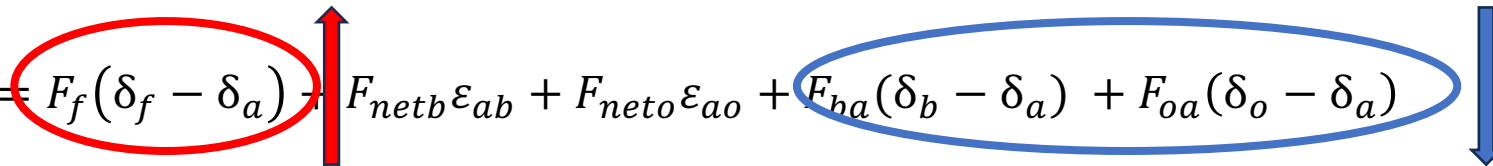
Midwest: LEF, WBI influenced by maize (C₄)
East: SCT, AMT influence more by C₃



Global mean $\delta^{13}\text{C}$ shows that much of CO_2 variability originates in the terrestrial biosphere.



Why is the $\delta^{13}\text{C}$ growth rate not trending?

$$C_a \frac{d\delta_a}{dt} = F_f(\delta_f - \delta_a) + F_{netb}\epsilon_{ab} + F_{neto}\epsilon_{ao} + F_{na}(\delta_b - \delta_a) + F_{oa}(\delta_o - \delta_a)$$


Over the last 30 years:

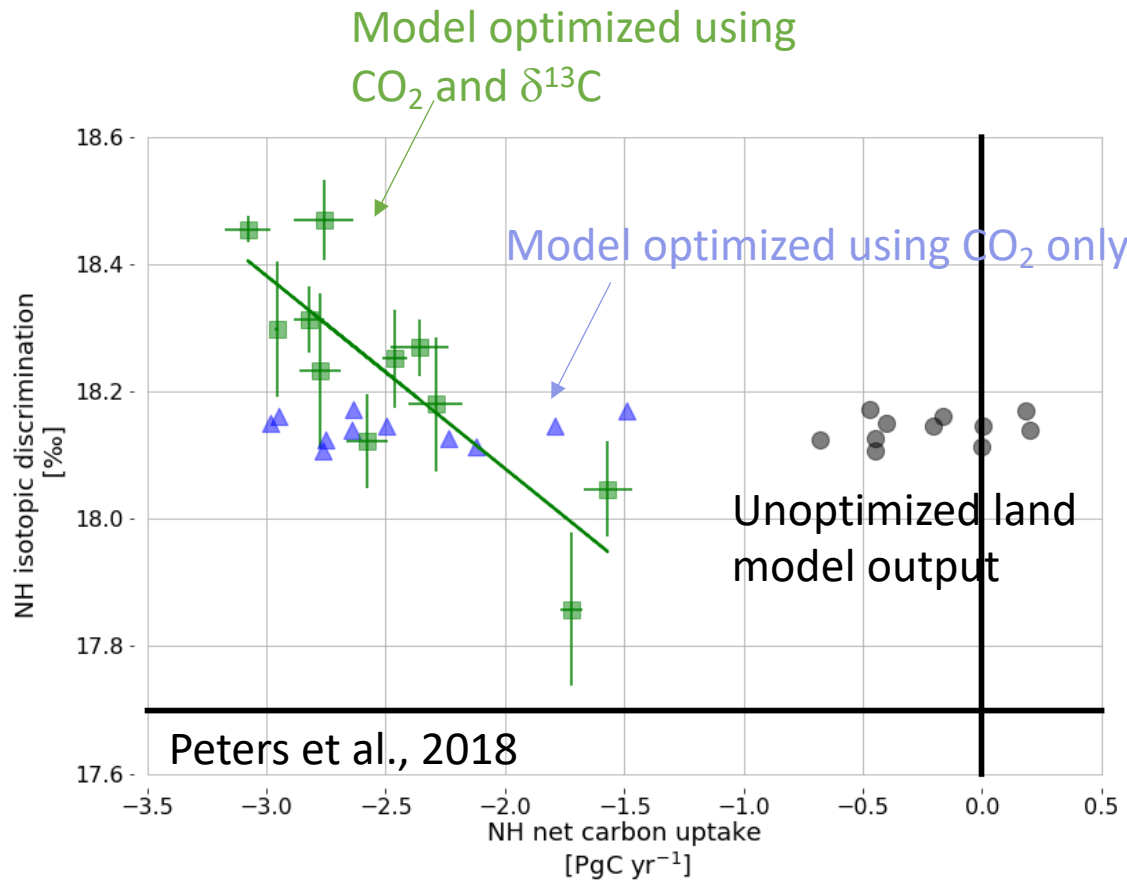
1. **Fossil isoflux** has decreased from ~ -120 to -200 $\text{PgC}^*(\text{per mil})/\text{yr}$
2. **Isotopic disequilibrium** has increased over the same period from ~ 80 to 140 $\text{PgC}^*(\text{per mil})/\text{yr}$
3. What does this mean?

$$\text{Diseqflux} \propto \text{GrossFlux} * \text{reservoir_residence_time}$$

→ Some combination of gross fluxes (e.g. respiration) and/or residence times in the surface waters or terrestrial biosphere have been changing a lot!
→ Our bottom-up calculations explain $\sim 85\%$ of this change (but our bottom-up absolute values are too low.)

$^{13}\text{CO}_2$ inversions can improve the drought response of land models.

Model response to drought



Recently developed SiB4 (land model)

