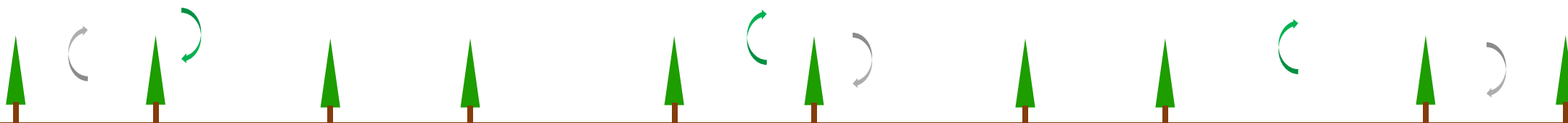


# Evaluation of the diurnal cycle of $O_2$ and $CO_2$ above the canopy of a forest and its application to further constrain the forest carbon balance

The different  $O_2:CO_2$  ratios above a forest and how to use them

**Kim Faassen**

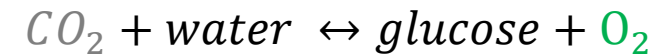
Linh Nguyen, Eadin Broekema, Bert Kers, Ivan Mammarella, Janne Levula, Timo Vesala, Penelope Pickers, Andrew Manning, Harro Meijer, Bert Heusinkveld, Wouter Peters, Jordi Vilá -Guerau de Arellano and Ingrid Luijkx



# The Exchange Ratio (ER)

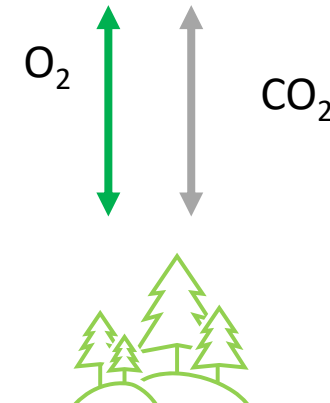
- Short time scale
- Continuous Time series
- Mole fractions
- Fluxes
- Local scale

$$ER = \frac{\Delta O_2}{\Delta CO_2}$$



$$ER/\alpha_B = -1.1$$

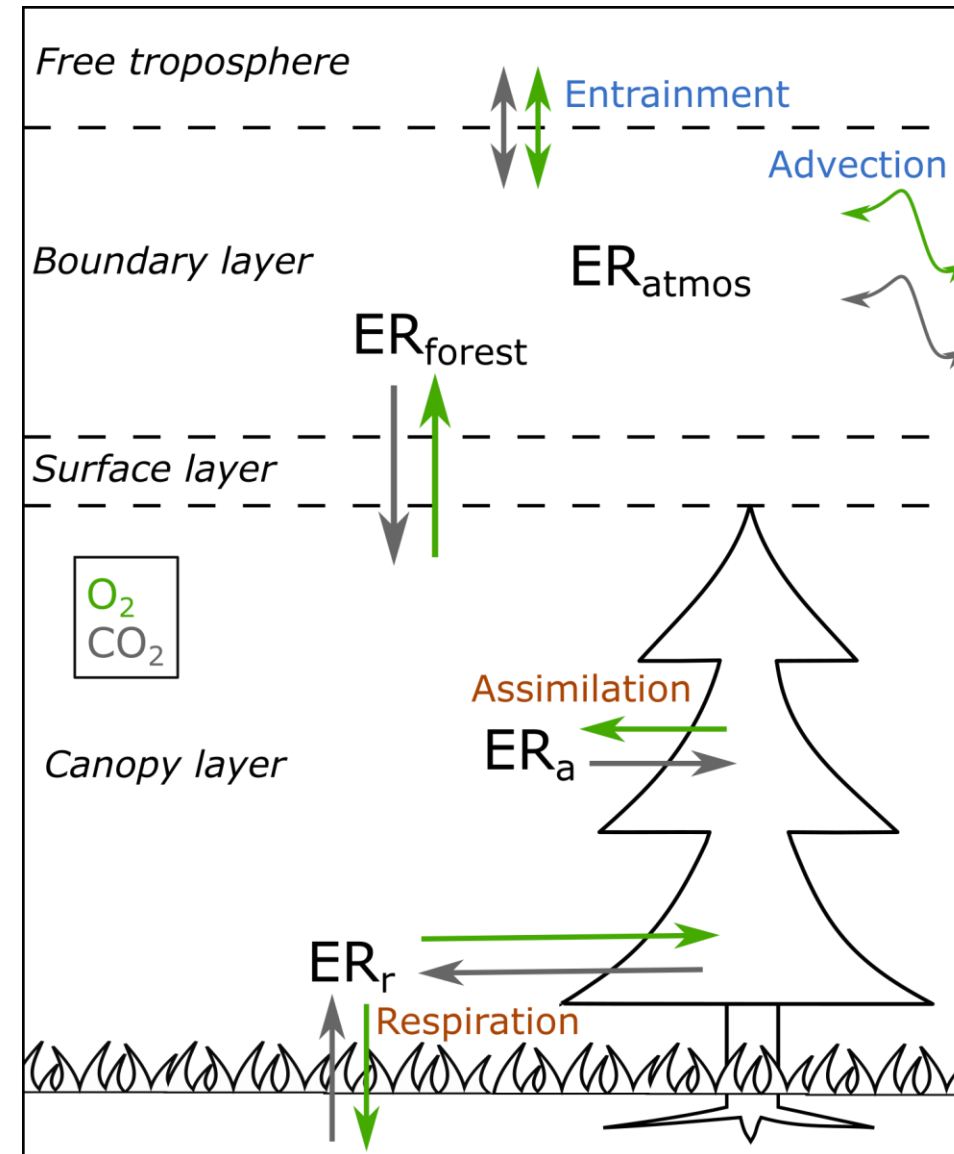
Severinghaus, 1995



# Local scale

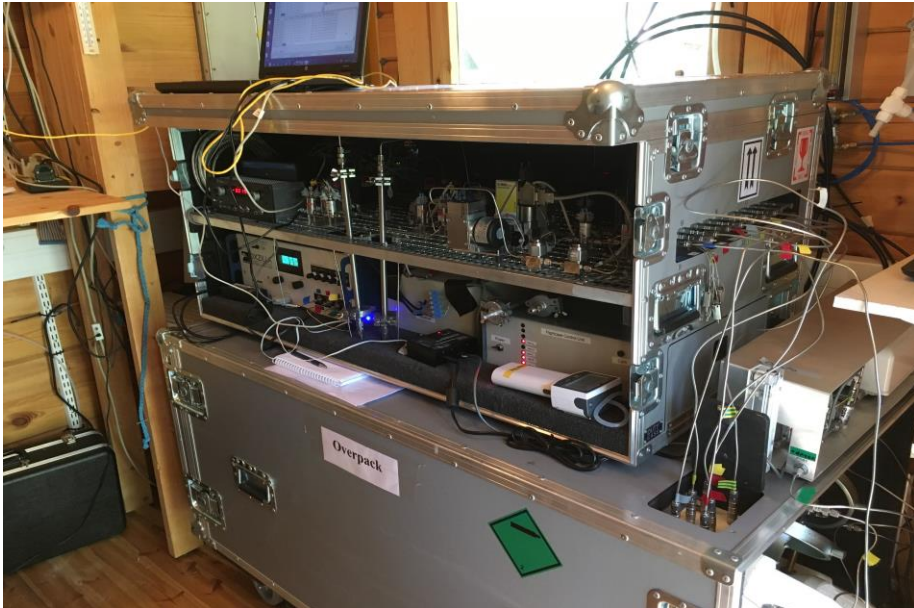
- $ER_{\text{atmos}}$  = 1 single height
- $ER_{\text{forest}}$  = surface fluxes
  - Multiple heights

*Get new insights into ER signals of boreal forest*

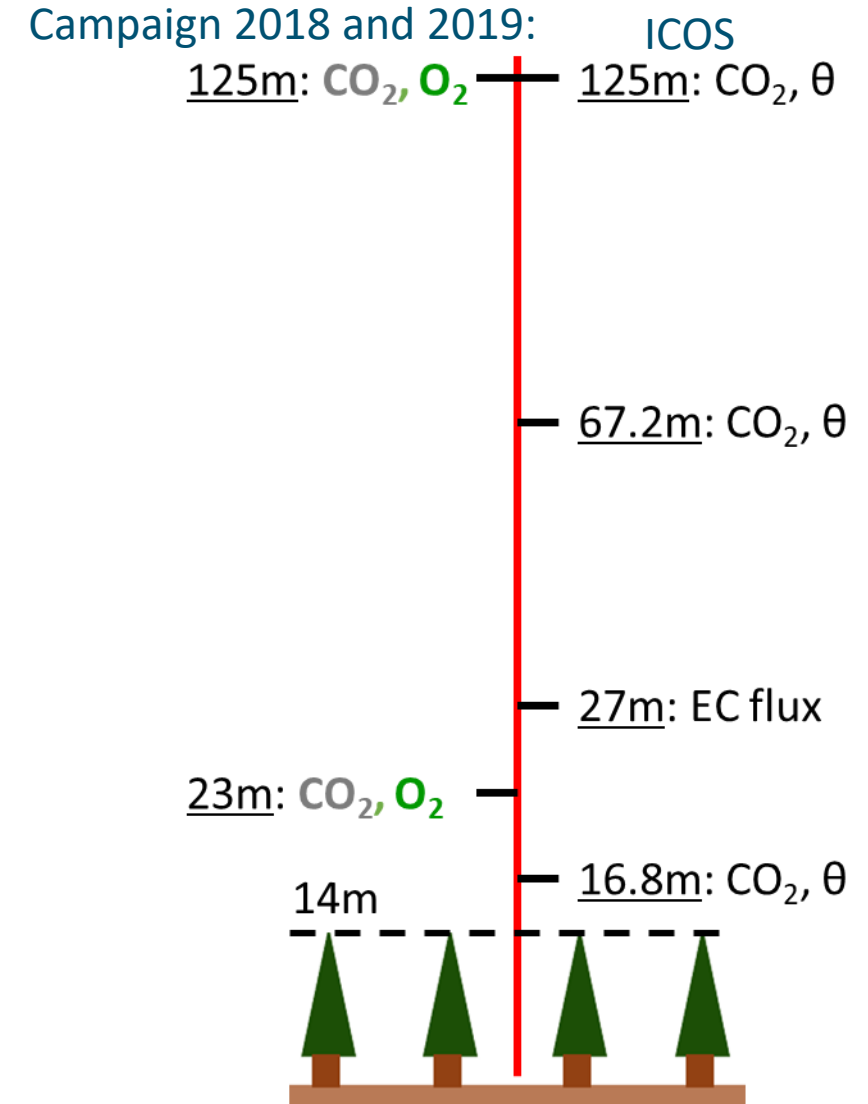


# Measurements at Hyytiälä

Oxzilla O<sub>2</sub> analyser & NDIR CO<sub>2</sub> analyser

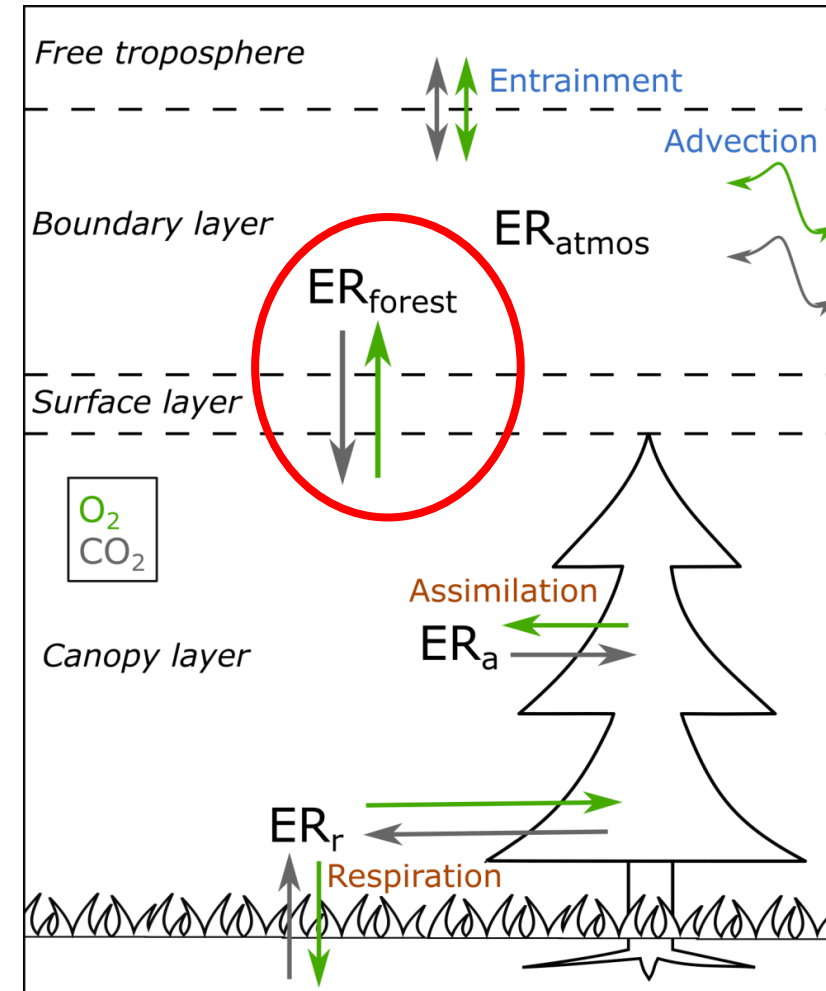


Faassen, K. A. P., Nguyen, L. N. T., Broekema, E. R., Kers, B. A. M., Mammarella, I., Vesala, T., Pickers, P.A., Manning, A.C., Vilà-Gueraude Arellano, J., Meijer, H.A.J., Peters, W., & Lujikx, I.T. (2023). Diurnal variability of atmospheric O<sub>2</sub>, CO<sub>2</sub>, and their exchange ratio above a boreal forest in southern Finland. *Atmospheric Chemistry and Physics*, 23(2), 851-876

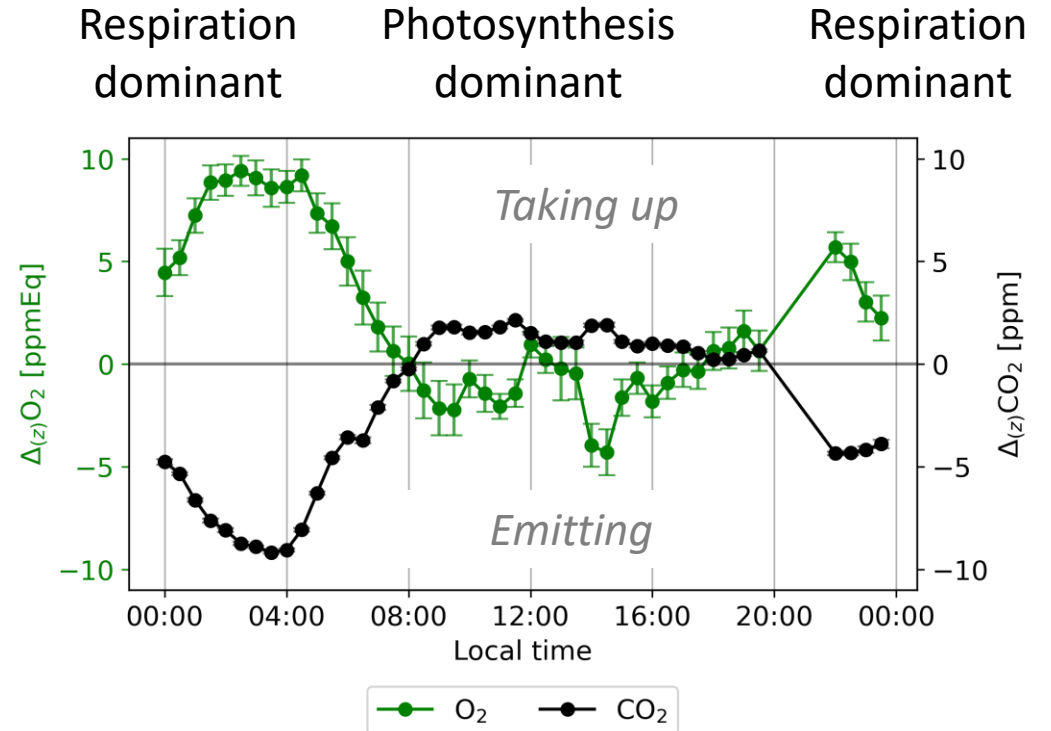
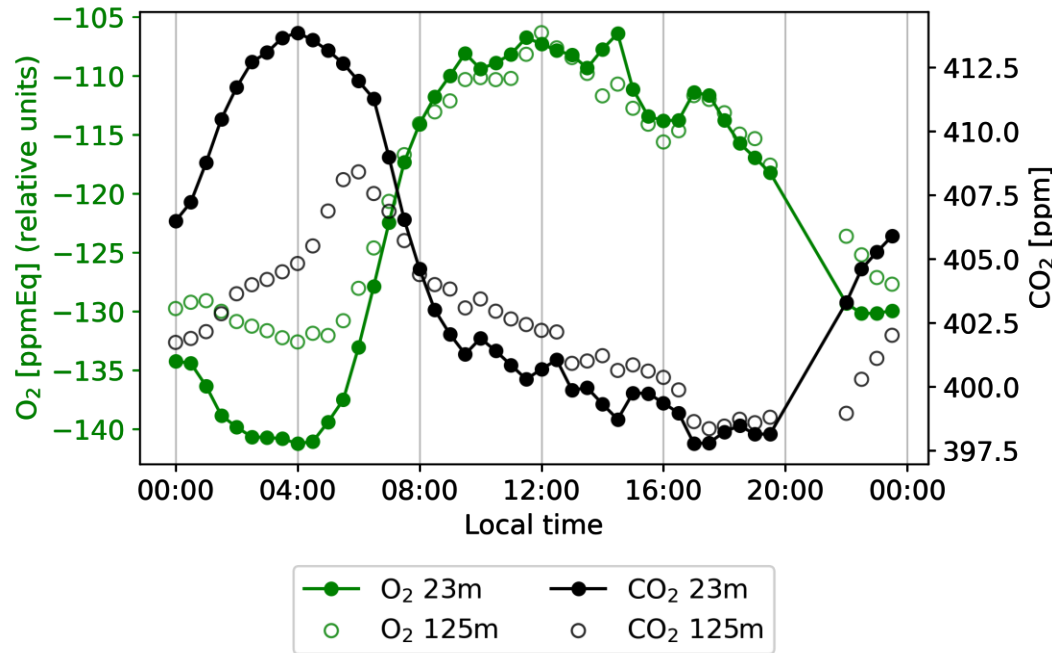


# $ER_{forest}$ : surface fluxes (multiple heights)

$$ER_{forest} = \frac{F(O_2)_s}{F(CO_2)_s}$$



# Gradient (125m – 23m):



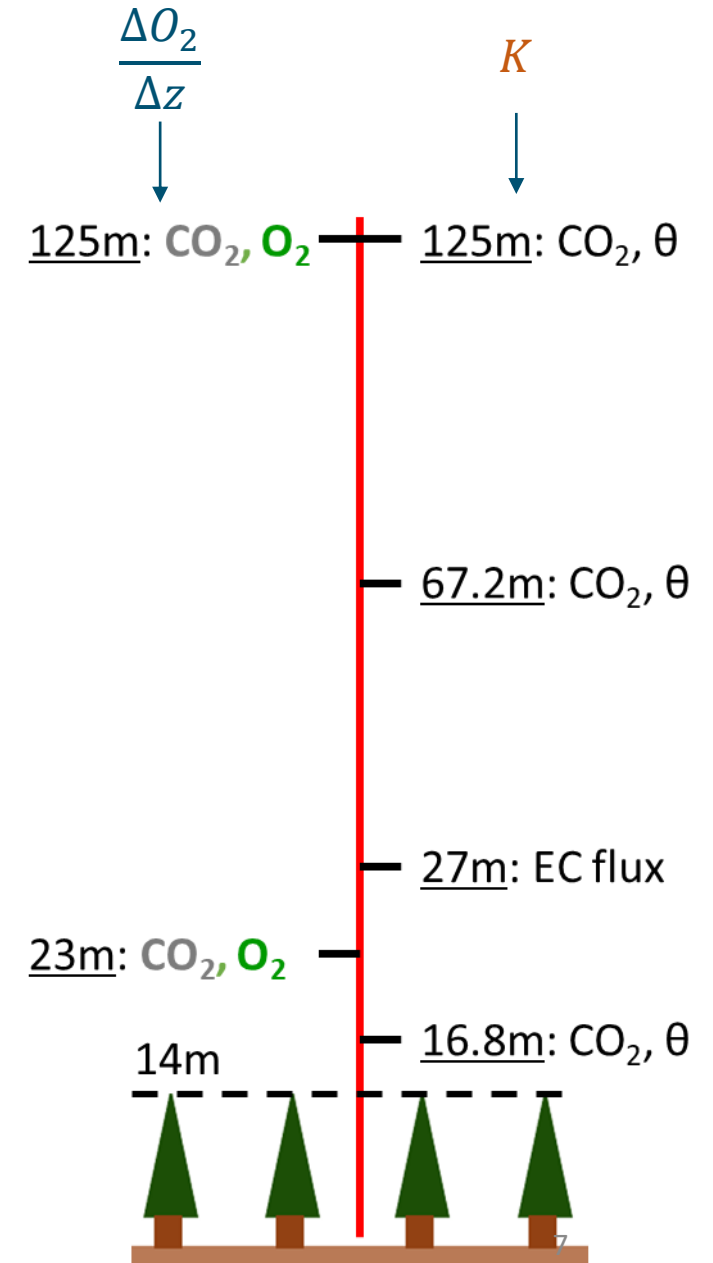
*Flux  $\approx$  -K\* Gradient*

# Flux calculations

*Flux  $\approx -K * \text{Gradient}$*

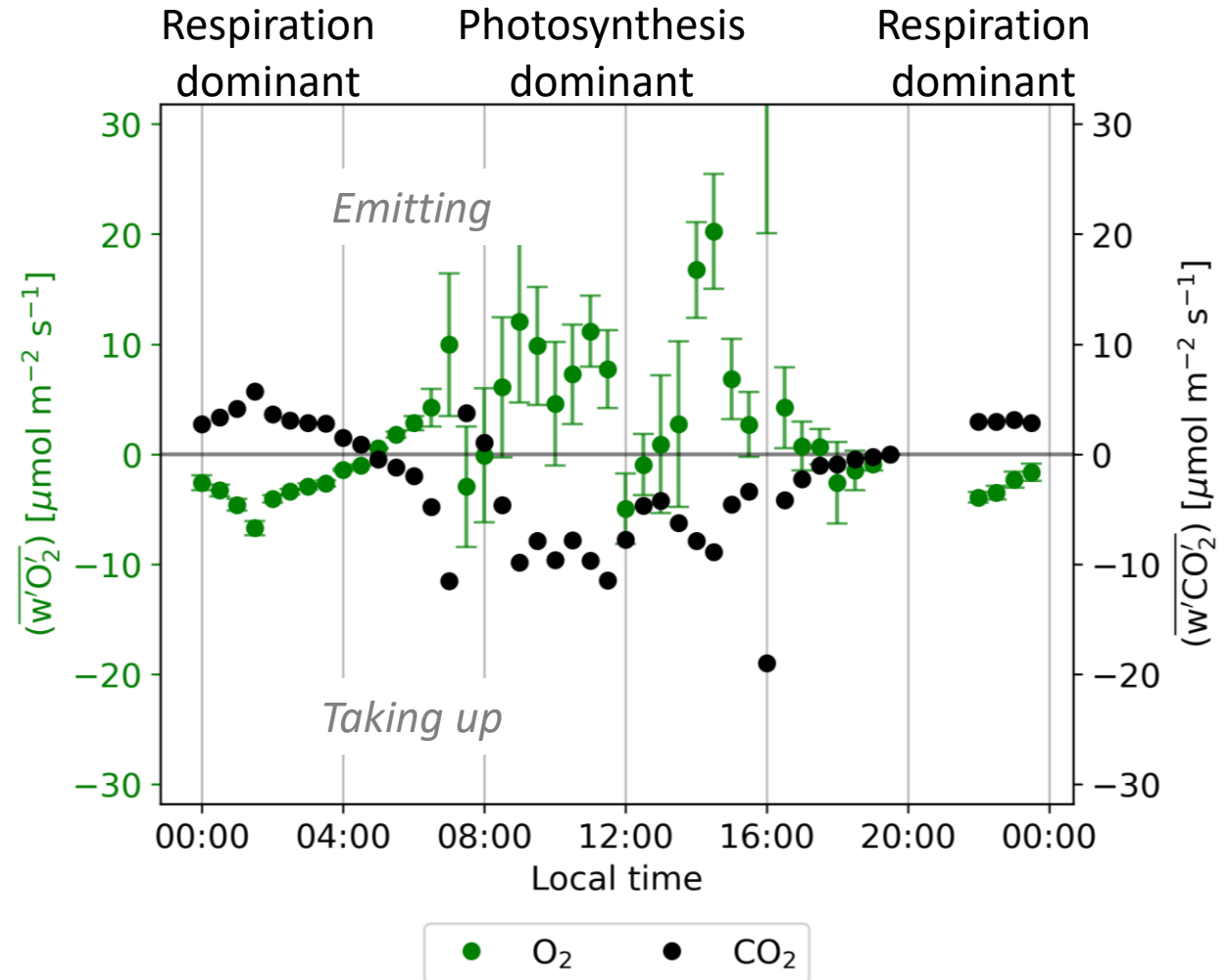
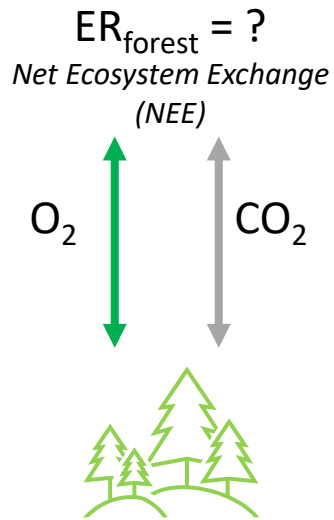
$$K = - \frac{\partial C O_2}{\partial z} / F(C O_2)_s$$

$$F(O_2)_s = -K * \frac{\Delta O_2}{\Delta z}$$



# O<sub>2</sub> and CO<sub>2</sub> surface fluxes

$Flux \approx -K * Gradient$



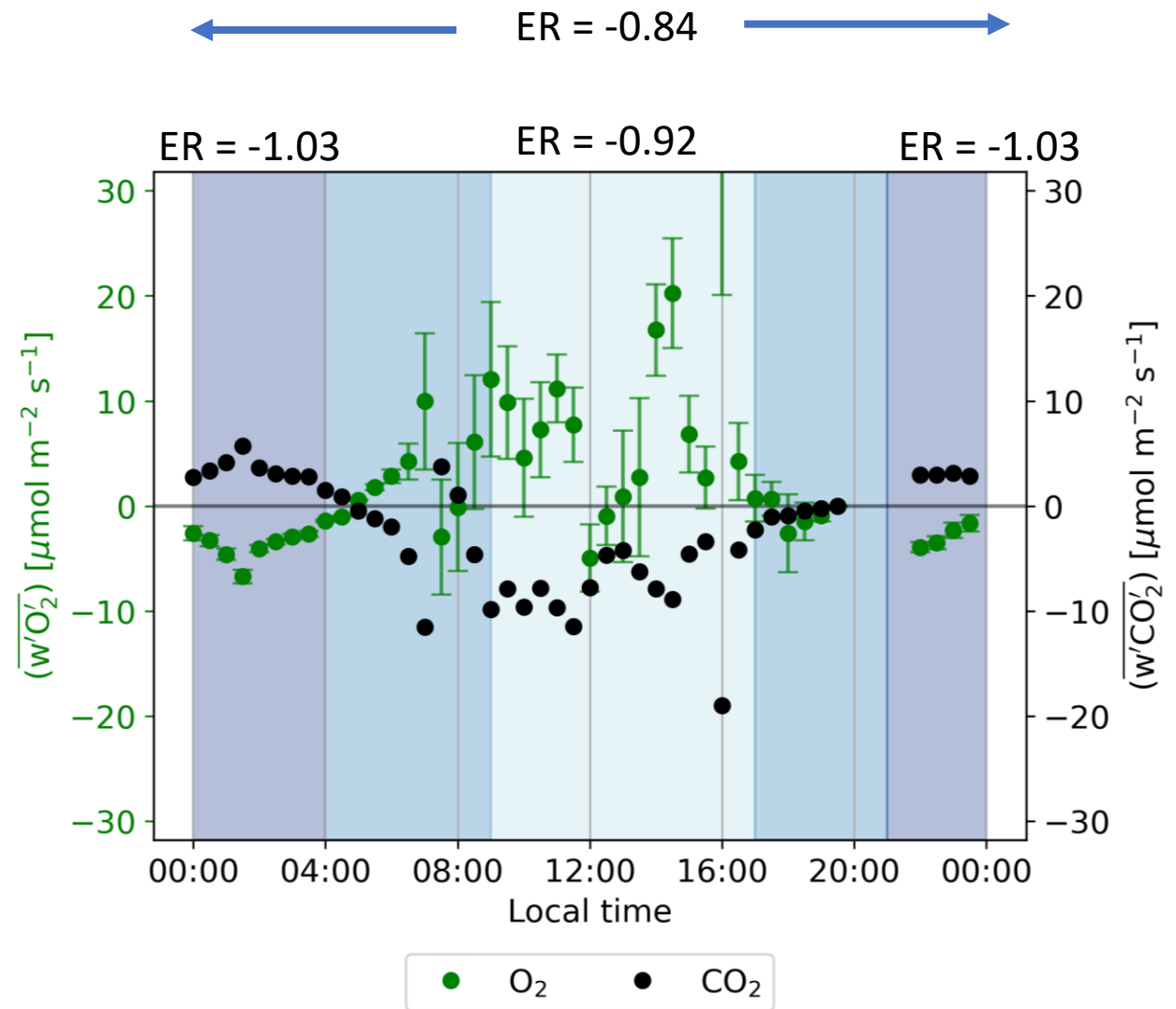
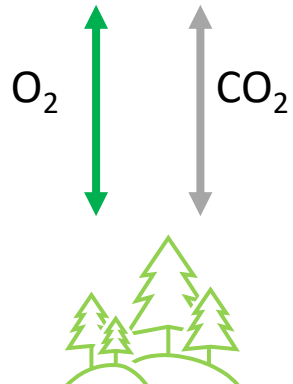
Faassen et al. (2023)



# ER<sub>forest</sub> signals

$$ER = \frac{\text{Surface flux } O_2}{\text{Surface flux } CO_2}$$

ER<sub>forest</sub> = -0.84  
Net Ecosystem Exchange (NEE)

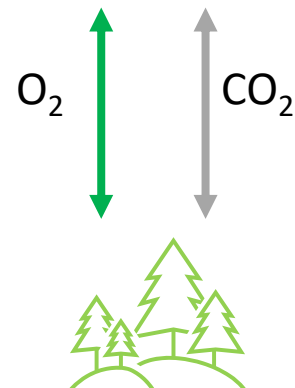


Faassen et al. (2023)

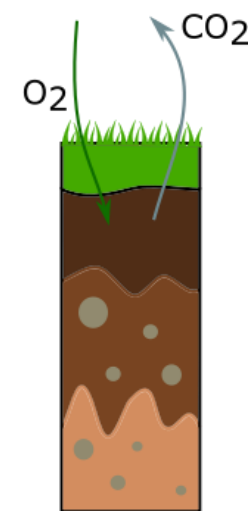
# Got values for ERa and ERr

- ER respiration = ER night
- ER assimilation = based on day ER and day GPP and TER

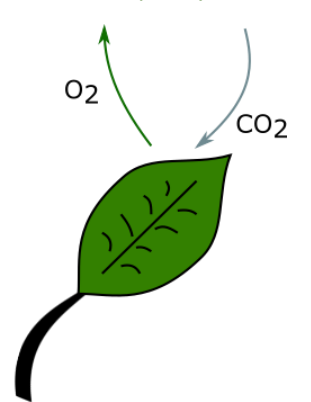
$ER_{\text{forest}} = -0.84$   
Net Ecosystem Exchange  
(NEE)



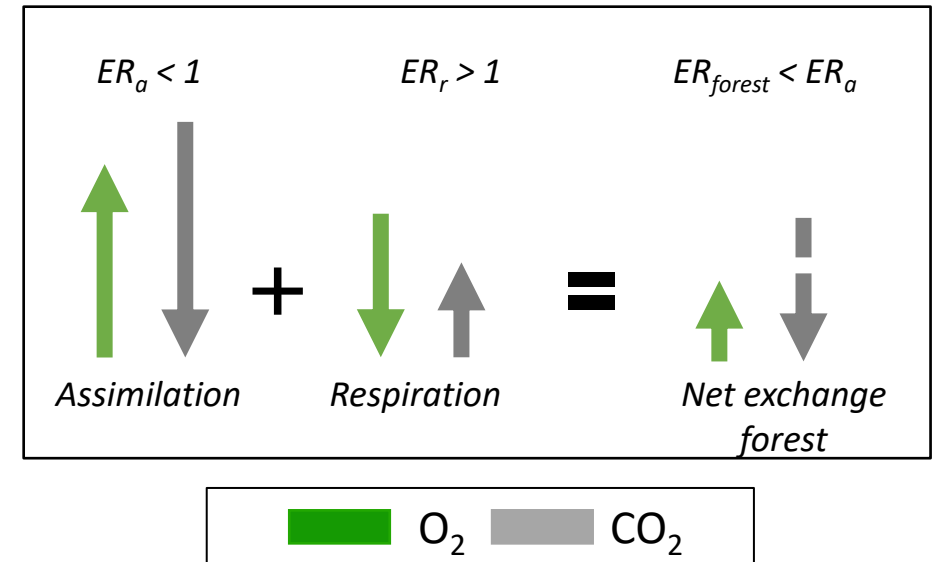
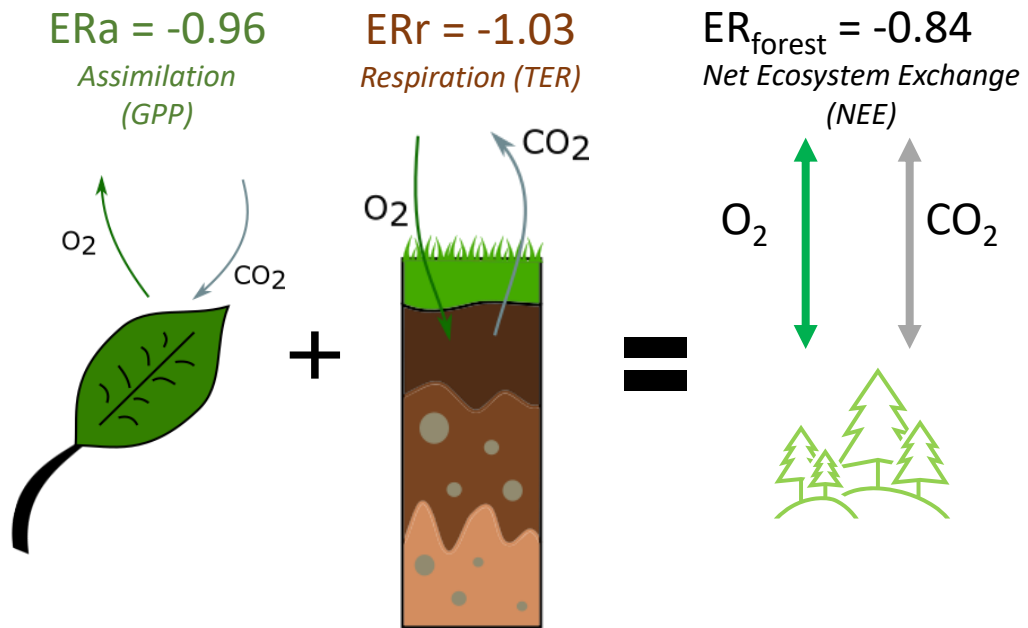
$ERr = -1.03$   
Respiration (TER)



$ERa = -0.96$   
Assimilation  
(GPP)



# Averaging of ER signals

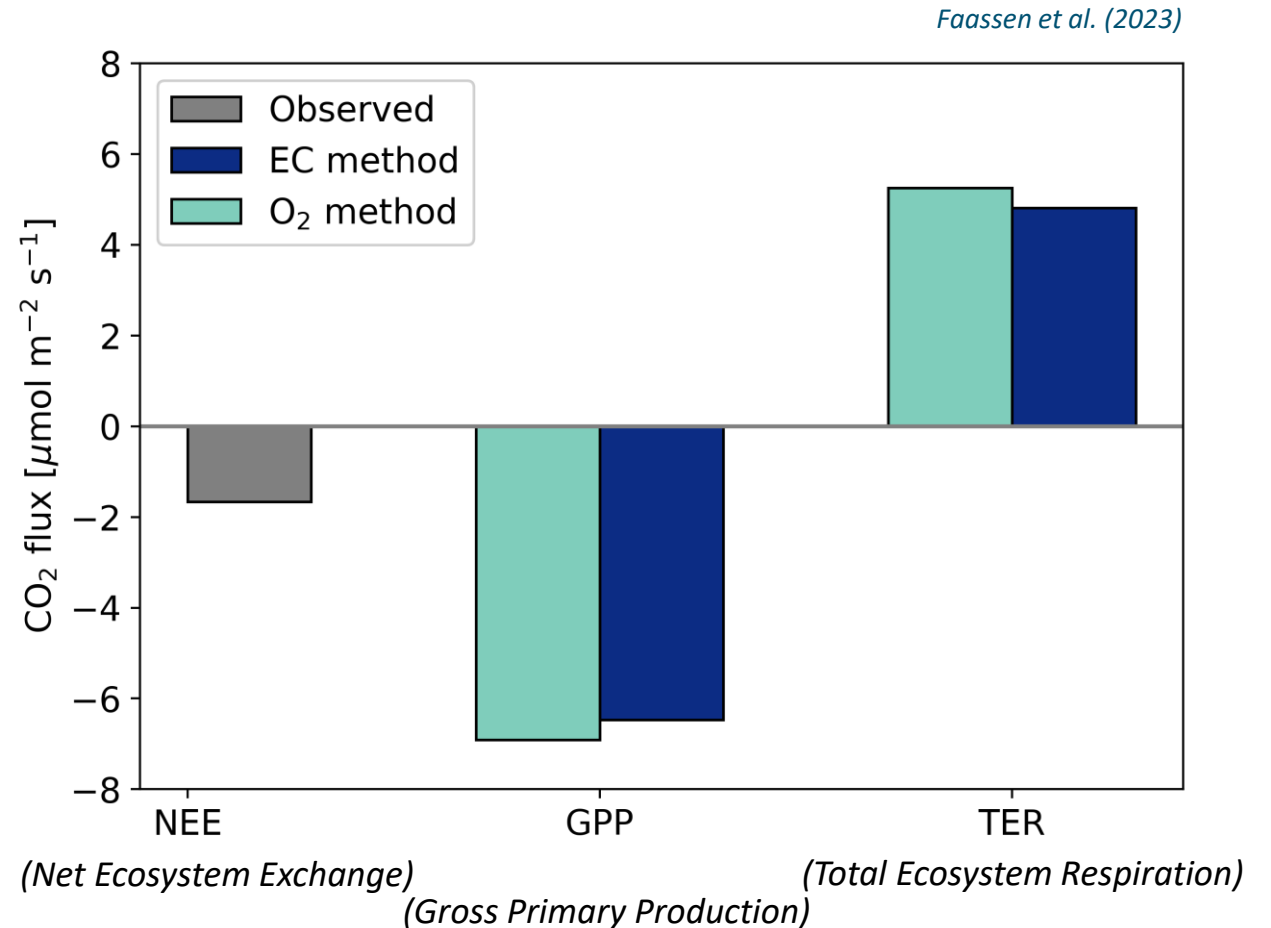


*ER values cannot be averaged when fluxes are in opposite directions.*

*Look at the  $O_2$  and  $CO_2$  budgets separately*

# O<sub>2</sub> as a partitioning method

- Applied to different day
- Shows promising results to use O<sub>2</sub> as a tool to split NEE into GPP and TER
- Works for 1 day in Hyytiälä



# New PhD project: O<sub>2</sub>/CO<sub>2</sub> measurements from lab to field



Lucas Hulsman



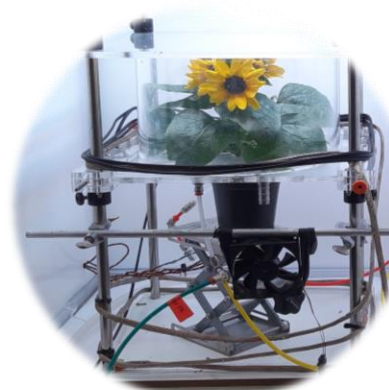
RQ: What is the variability of the ER on leaf- plant- and ecosystem- scale under different environmental conditions?

Ecosystem

Leaf

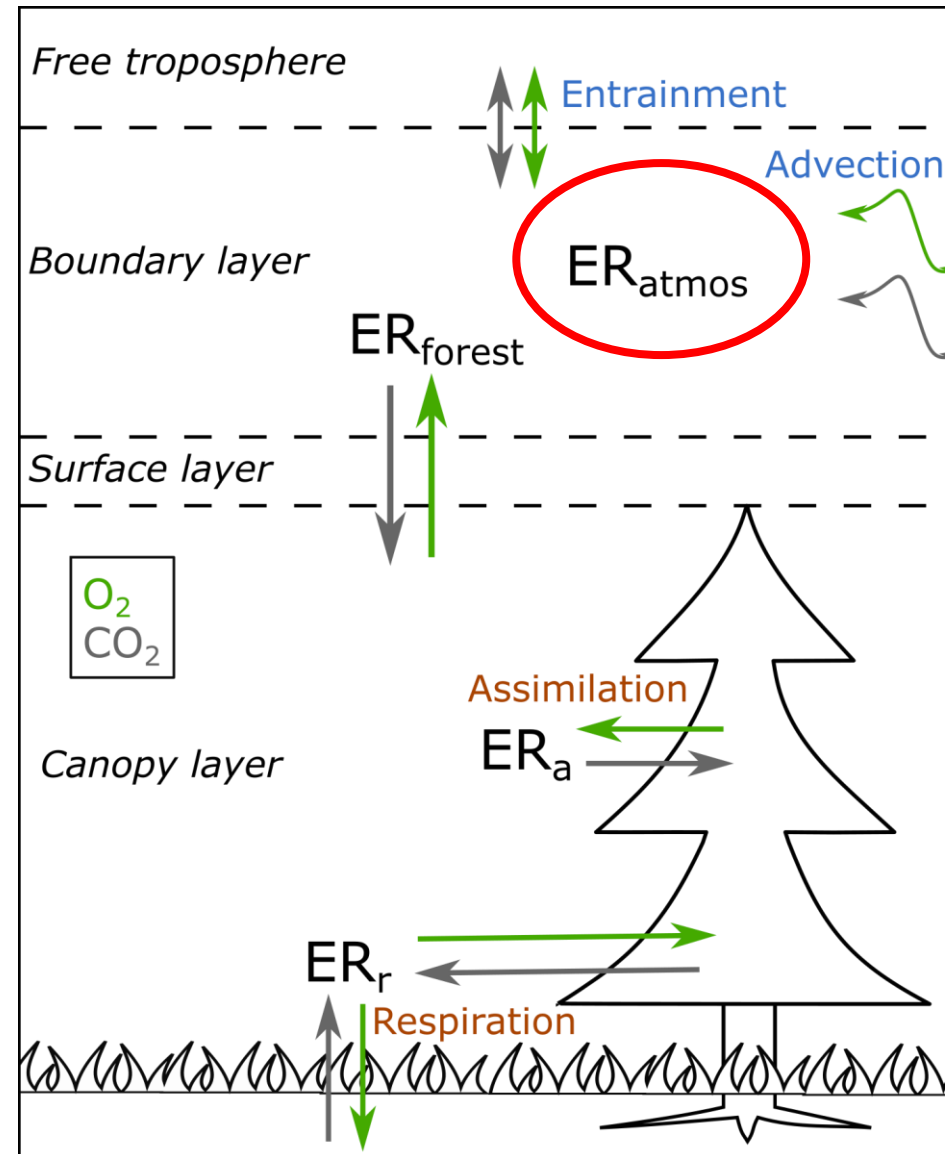


Plant

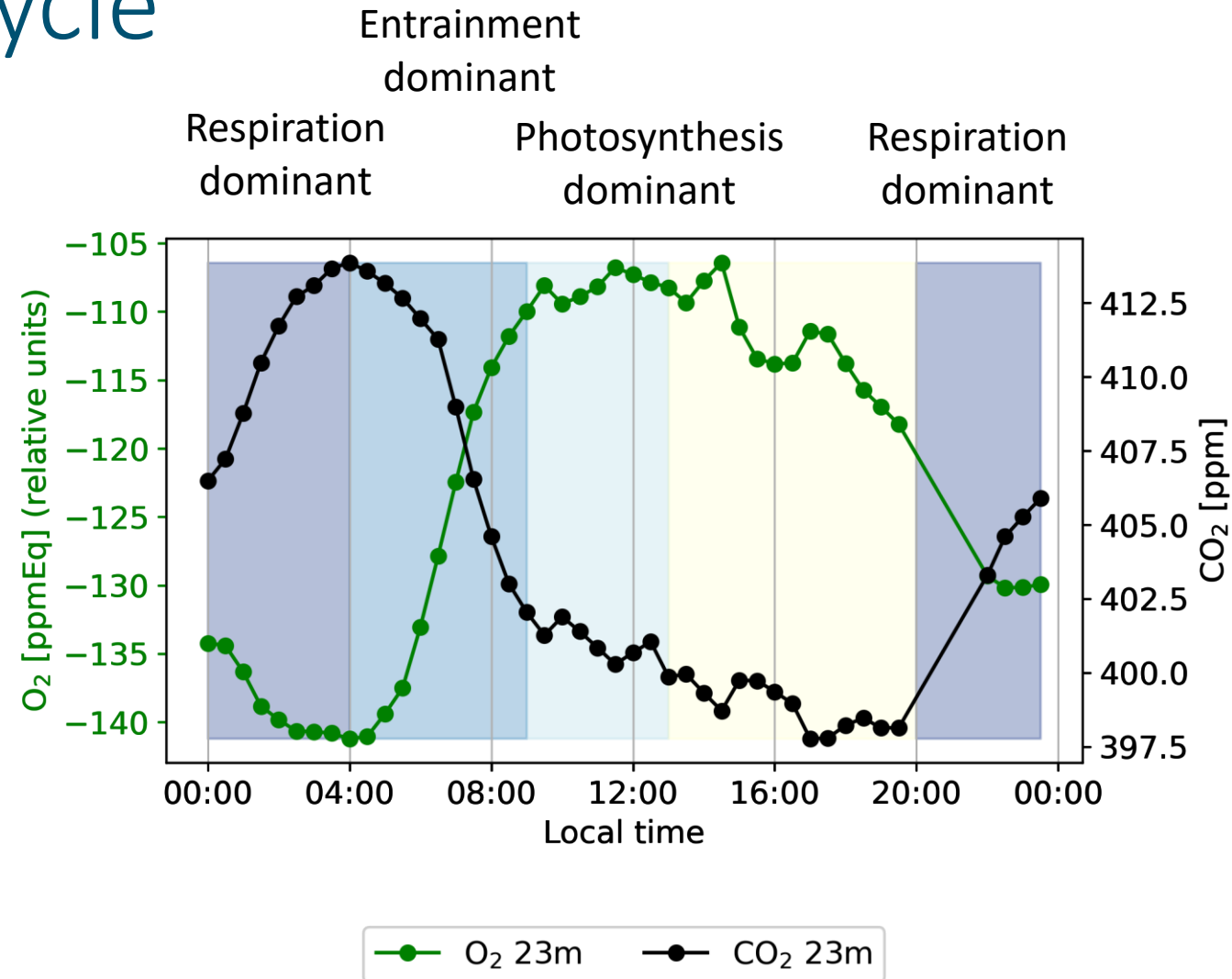


# $ER_{atmos}$ : 1 height

$$ER_{atmos} = \frac{\Delta(t)O_2}{\Delta(t)CO_2}$$



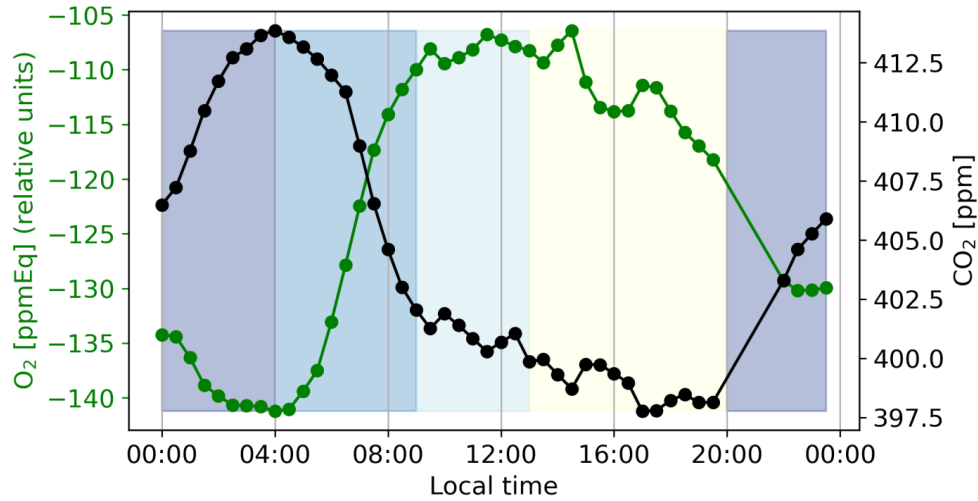
# Diurnal cycle



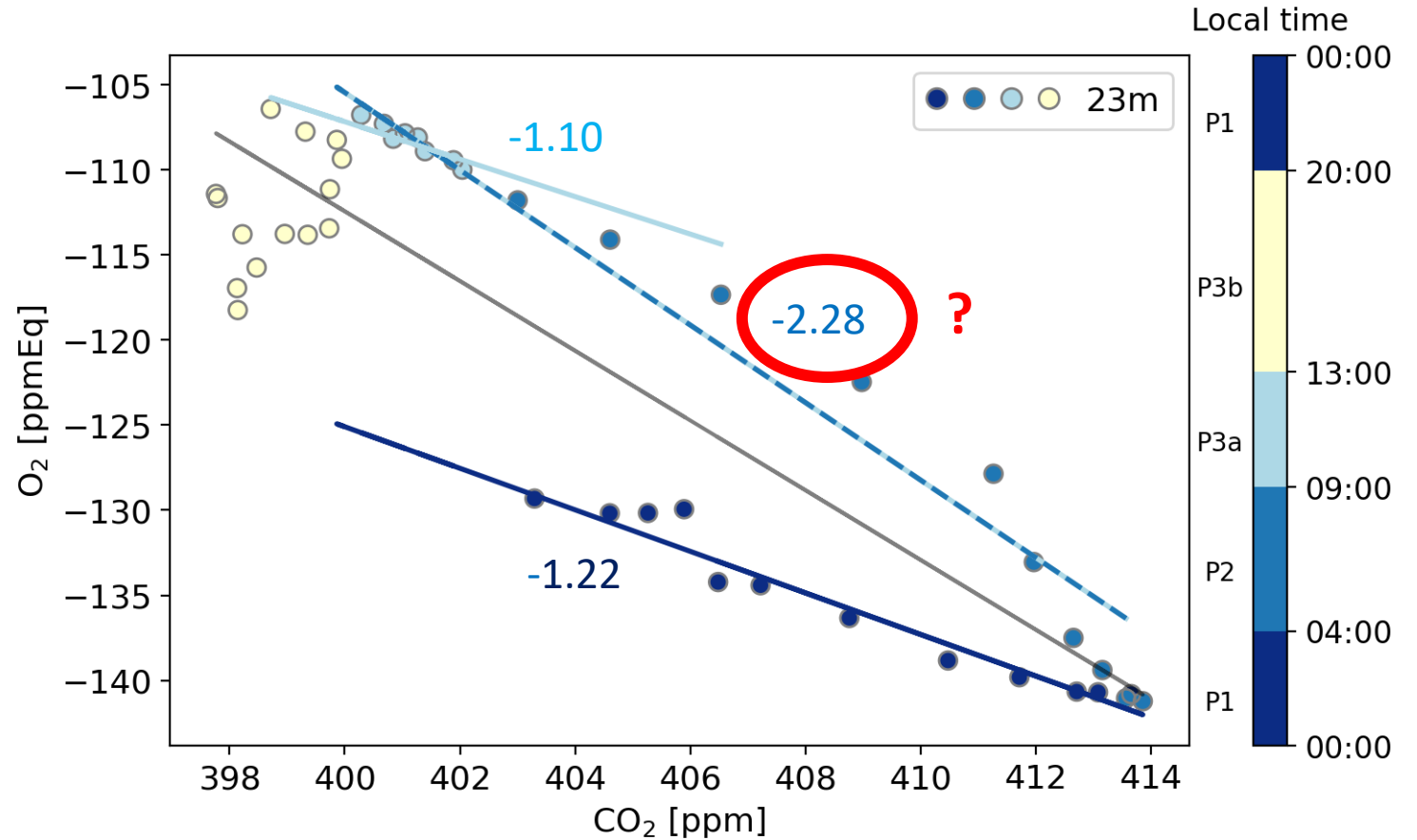
*Faassen et al. (2023)*

# ER<sub>atmos</sub> signals

$$ER = \frac{\Delta(t)O_2}{\Delta(t)CO_2}$$



—●— O<sub>2</sub> 23m    —●— CO<sub>2</sub> 23m



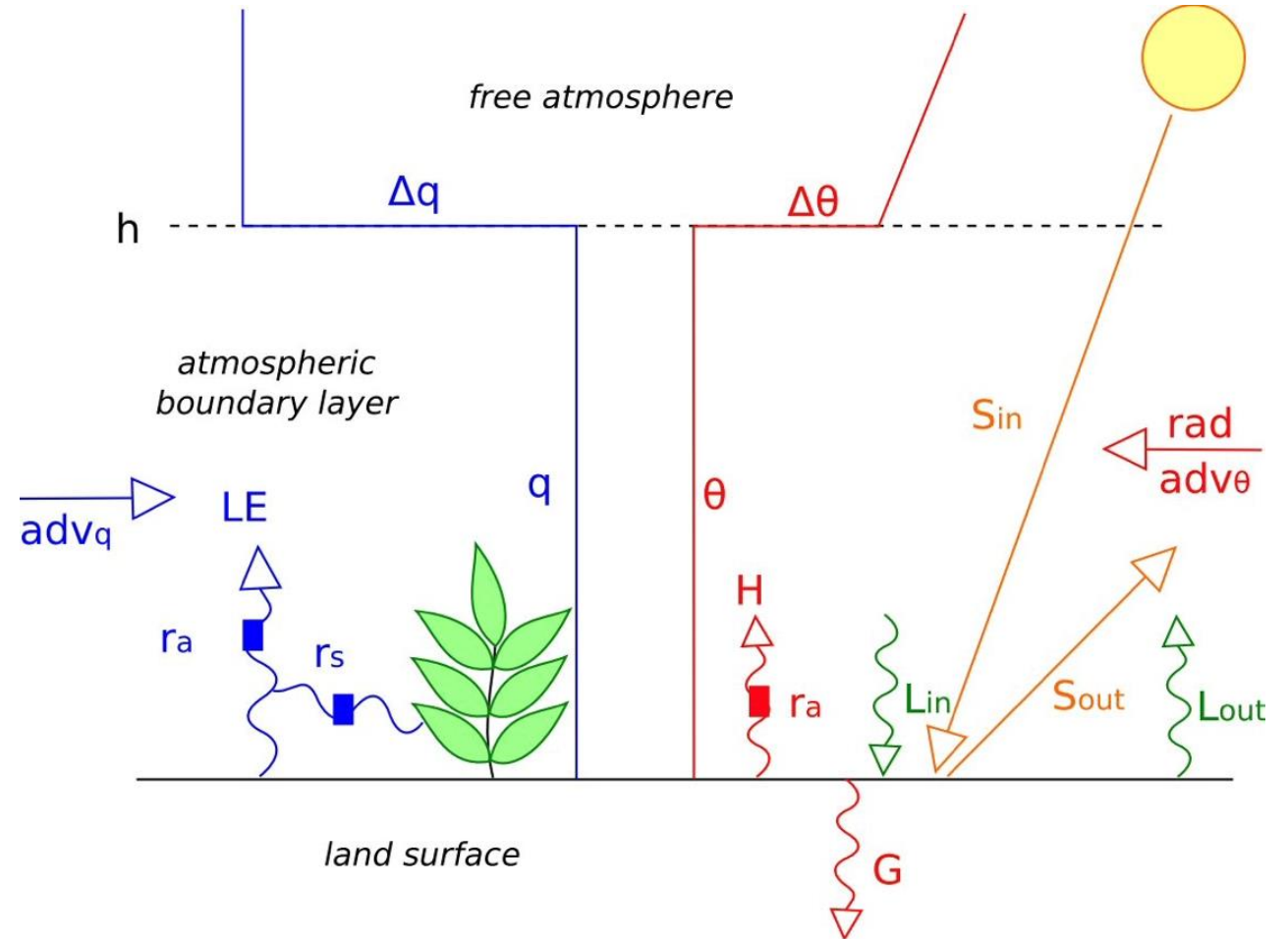
*Faassen et al. (2023)*

*Explain ER<sub>atmos</sub> signal*



# CLASS

- Chemistry Land-surface  
Atmosphere Soil Slab model
- 1-Box model (growing)
- Assumes mixed-layer theory
- Runs for 1 day
- Added  $O_2$  (Faassen et al. (in prep))



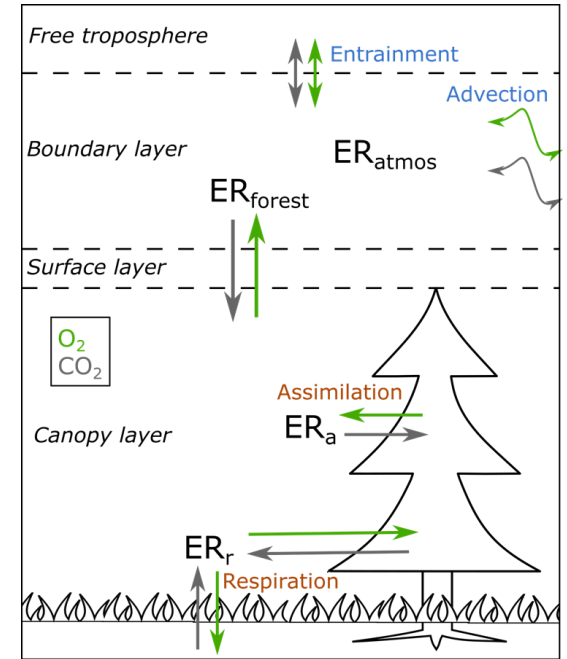
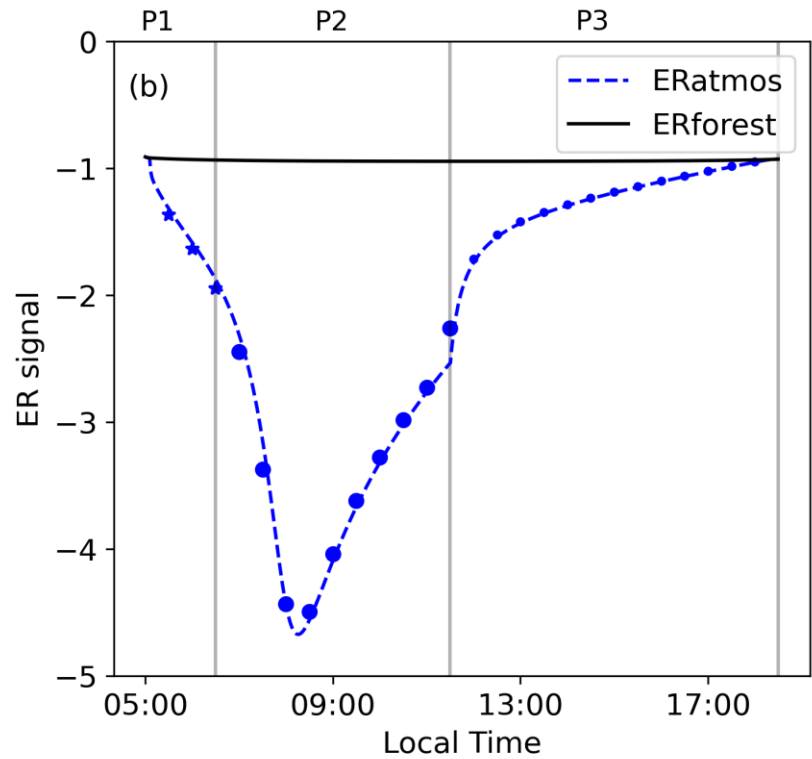
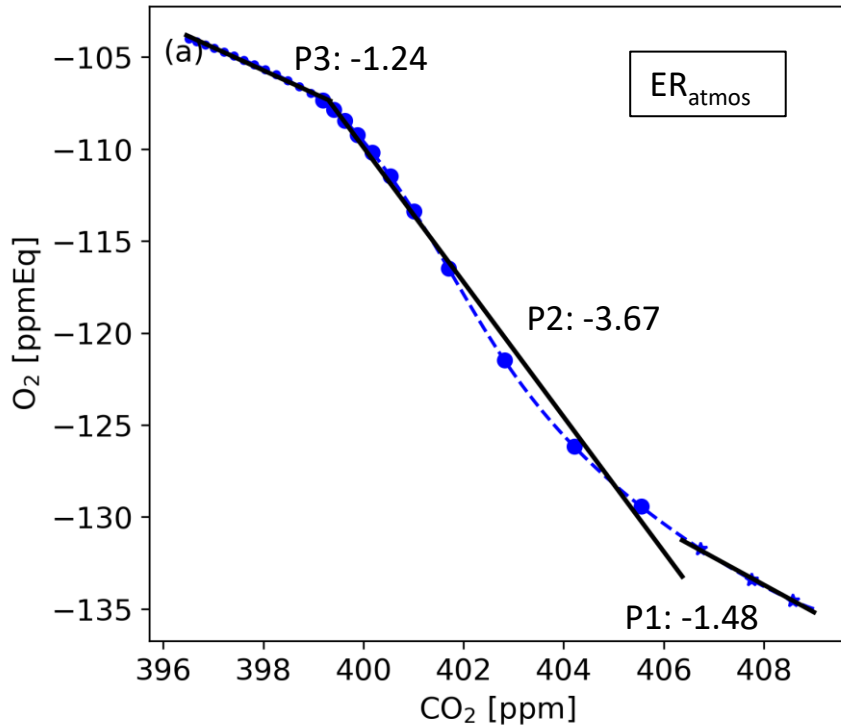
*Jacobs & de Bruin 1997; Ronda et al., 2001; Jacobs et al., 2007*

Vilá-Guerau de Arellano, J., van Heerwaarden, C. C., van Stratum, B. J. H., and van den Dries, K.:  
The Atmospheric Boundary Layer, Cambridge University Press, 2015

# Model O<sub>2</sub> with CLASS

$$ER_{atmos} = \frac{\Delta(t)}{\Delta(t)}$$

$$ER_{forest} = \frac{F(\Delta(z))}{F(\Delta(z))}$$

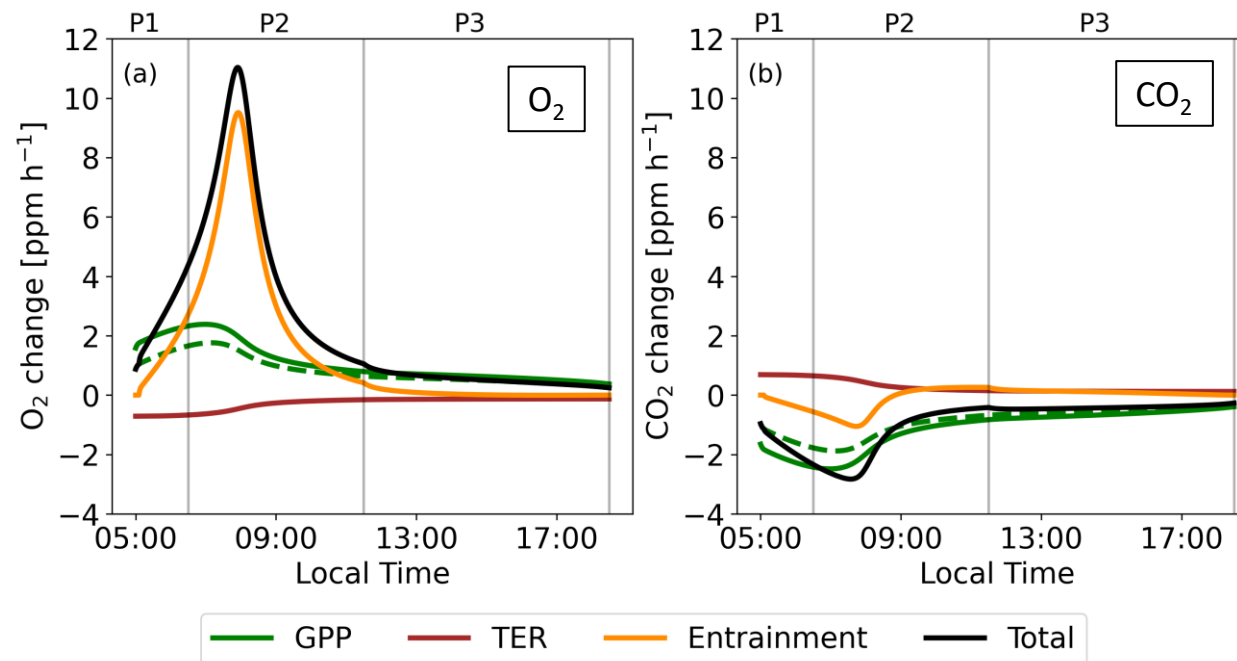
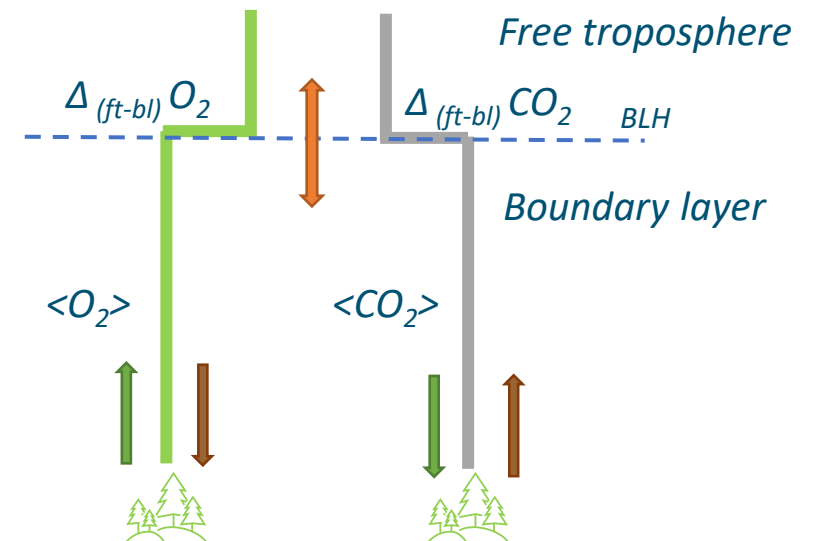


We can reproduce the high values for ER<sub>atmos</sub> that we observed in Hyytiälä

Explain P2: why such high values?  
Explain P3: can ER<sub>atmos</sub> be used as ER<sub>forest</sub>

# Importance of entrainment

- First clue look at tendencies
  - Entrainment is important
  - Both budgets behave differently



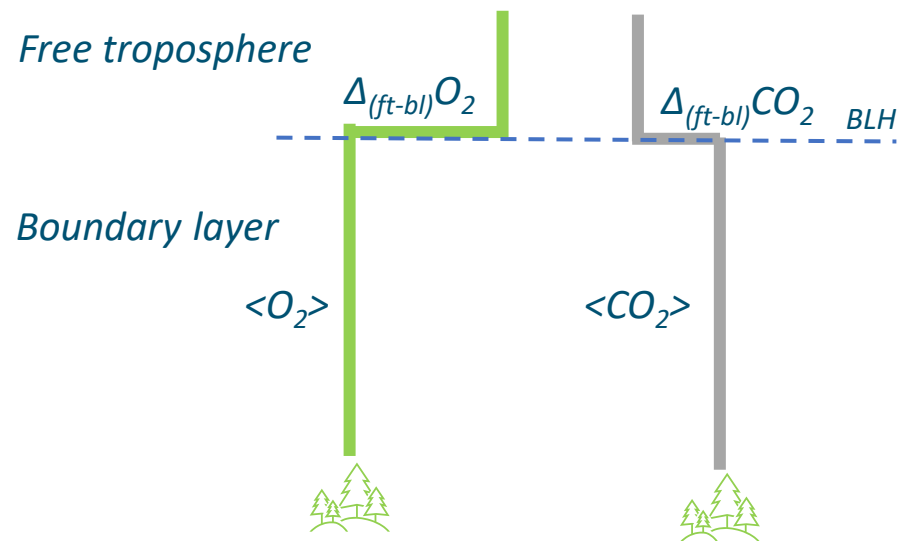
# Relationship between $ER_{forest}$ and $ER_{atmos}$

$$ER_{forest} = \frac{(F_{O_2})_s}{(F_{CO_2})_s}$$

$$ER_{atmos} = \frac{dO_2/dt}{dCO_2/dt} = \frac{((F_{O_2})_s - (F_{O_2})_e)/h}{((F_{CO_2})_s - (F_{CO_2})_e)/h}$$

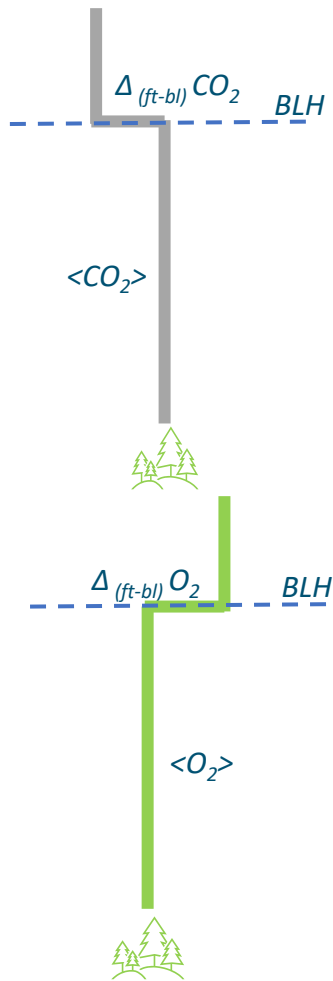
$$ER_{atmos} = ER_{forest} * \left( \frac{1 + \beta_{O_2}}{1 + \beta_{CO_2}} \right)$$

$$\left( \frac{1 + \beta_{O_2}}{1 + \beta_{CO_2}} \right) = \frac{1 + (w_e * \Delta_{(ft-bl)O_2}) / (F_{O_2})_s}{1 + (w_e * \Delta_{(ft-bl)CO_2}) / (F_{CO_2})_s}$$



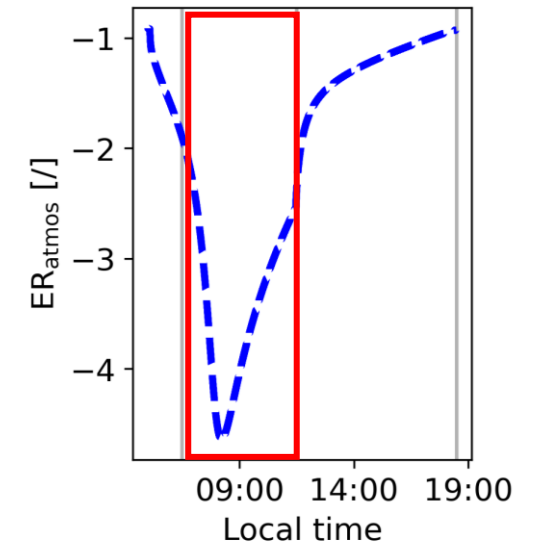
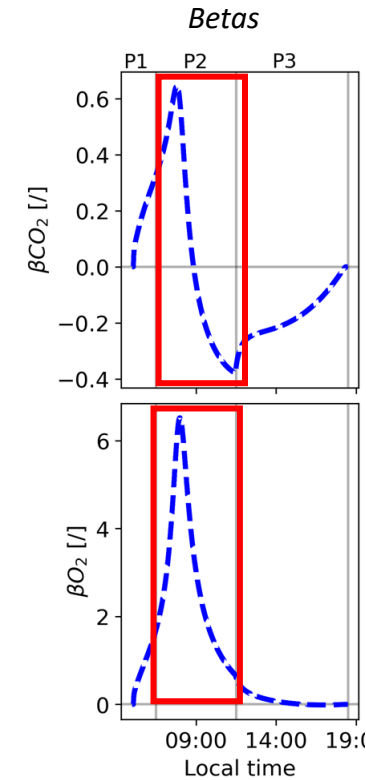
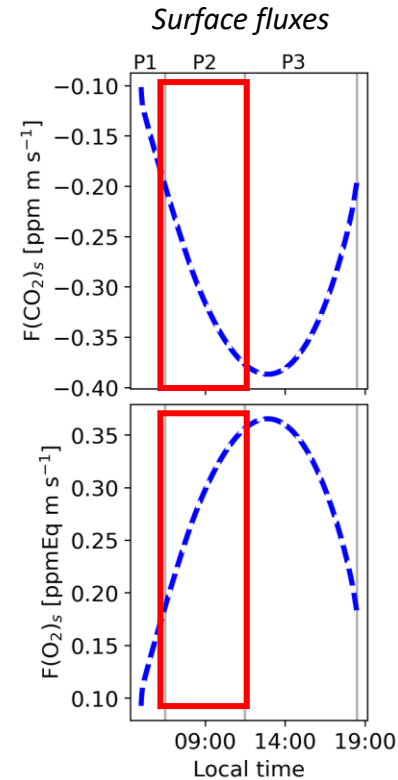
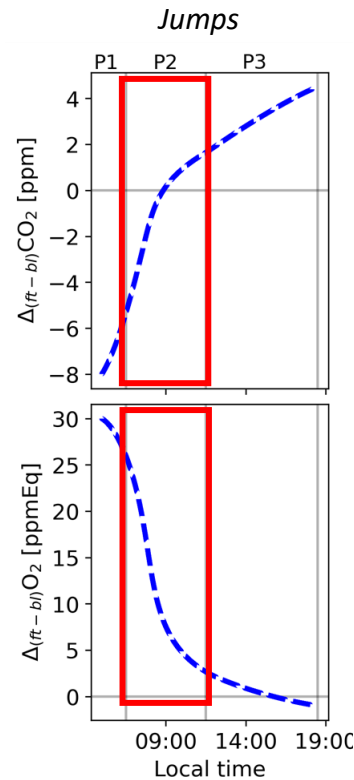
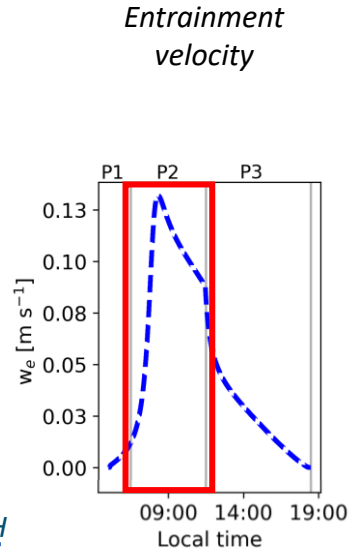
- Entrainment
- Jumps
- Surface fluxes

# Disentangle the equation: P2 (morning)



$$ER_{atmos} = ER_{forest} * \left( \frac{1 + \beta_{O_2}}{1 + \beta_{CO_2}} \right)$$

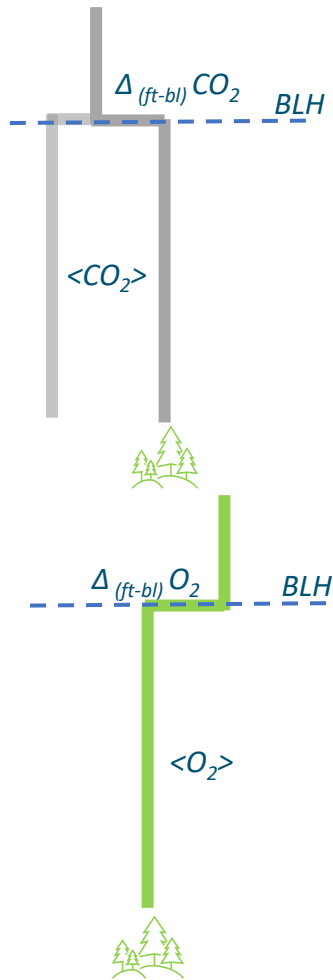
$$\beta = \frac{F_e}{F_s} = \frac{w_e * \Delta(ft-bl)}{F_s}$$



- Jump size  $\rightarrow \Delta_{(ft-bl)}$
- ERforest  $\rightarrow FO_2 : FCO_2$
- Jump ratio  $\rightarrow \Delta O_2 : \Delta CO_2$

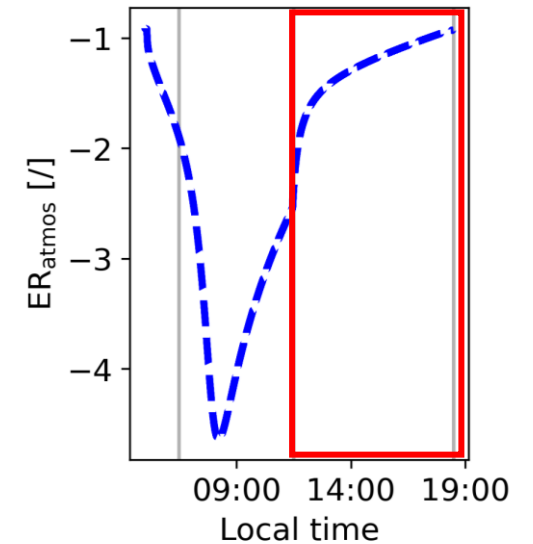
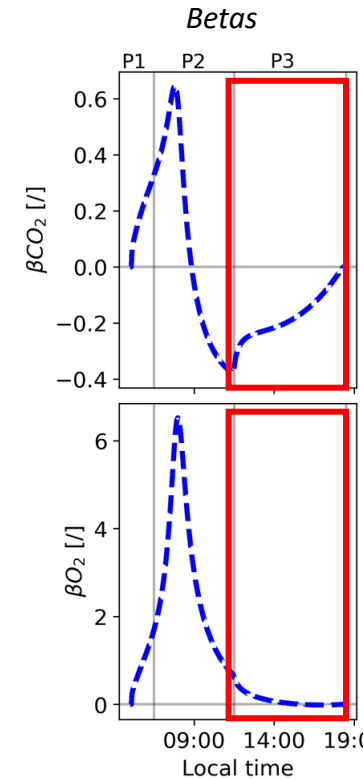
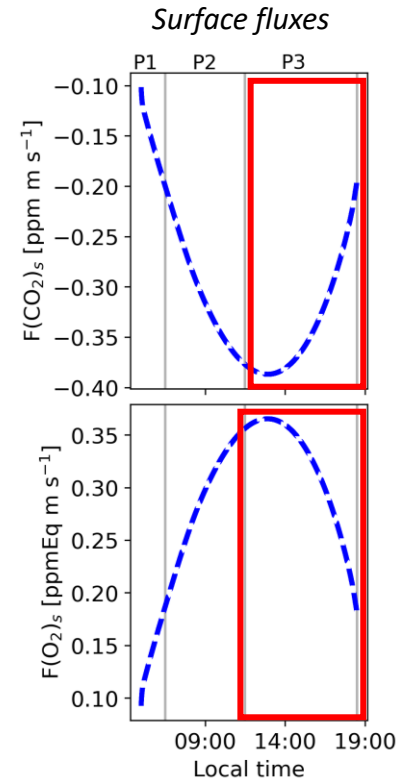
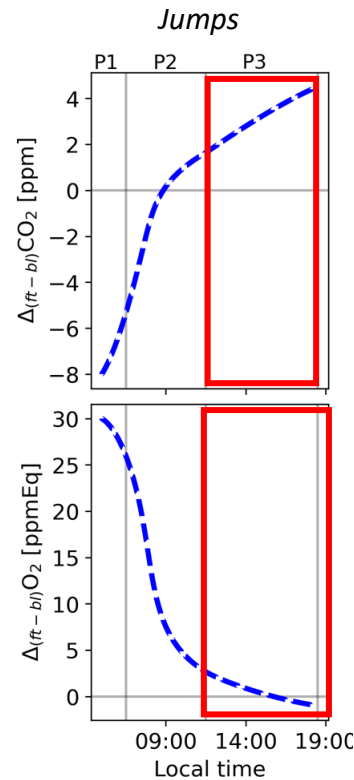
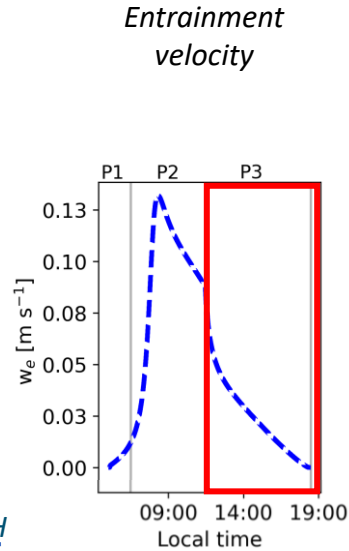
Could  $ER_{atmos}$  be used for  $ER_{forest}$ ?

# Disentangle the equation: P3 (afternoon)



$$ER_{atmos} = ER_{forest} * \left( \frac{1 + \beta_{O_2}}{1 + \beta_{CO_2}} \right)$$

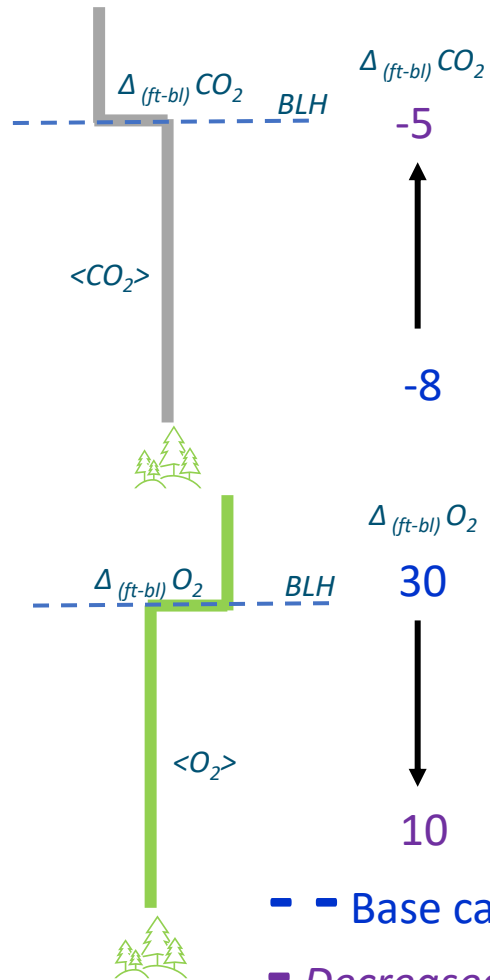
$$\beta = \frac{F_e}{F_s} = \frac{w_e * \Delta(ft-bl)}{F_s}$$



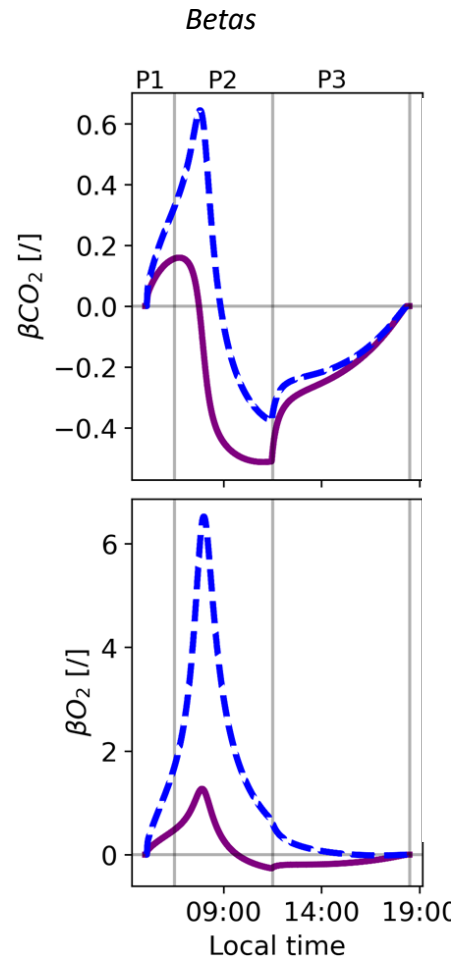
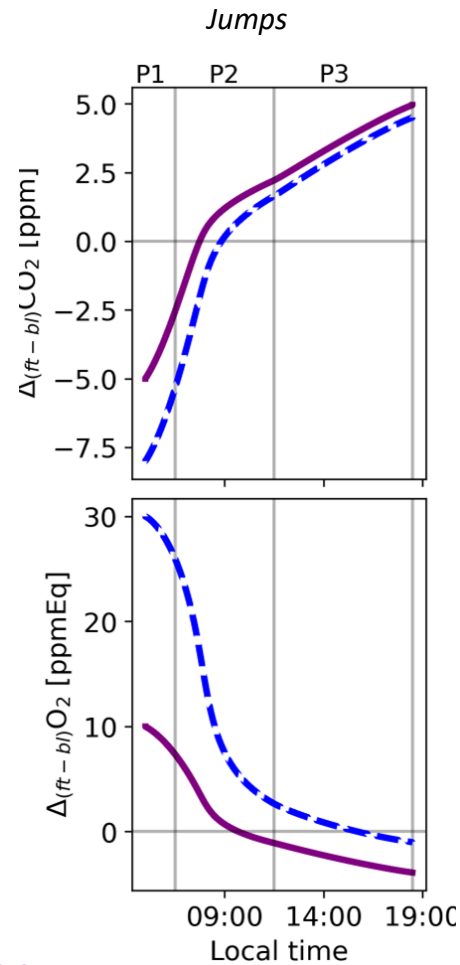
- Jump size  $\rightarrow \Delta_{(ft-bl)}$
- Flux size  $\rightarrow F(\Delta_{(z)})$
- Jump ratio  $\rightarrow \Delta O_2 : \Delta CO_2$

*Is this a specific case?*

# Other option: decrease jump size ( $\Delta_{(ft-bl)}$ )

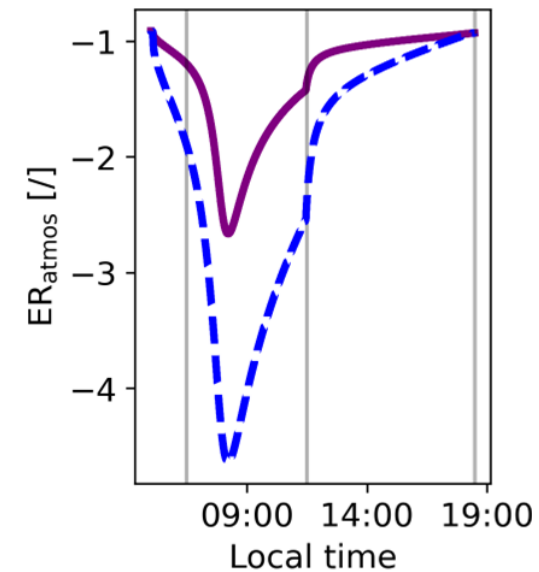


- - Base case  
 - Decreased jump



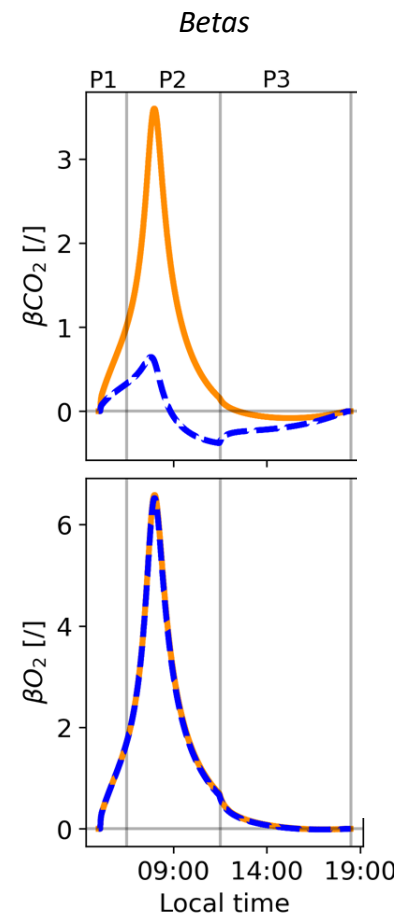
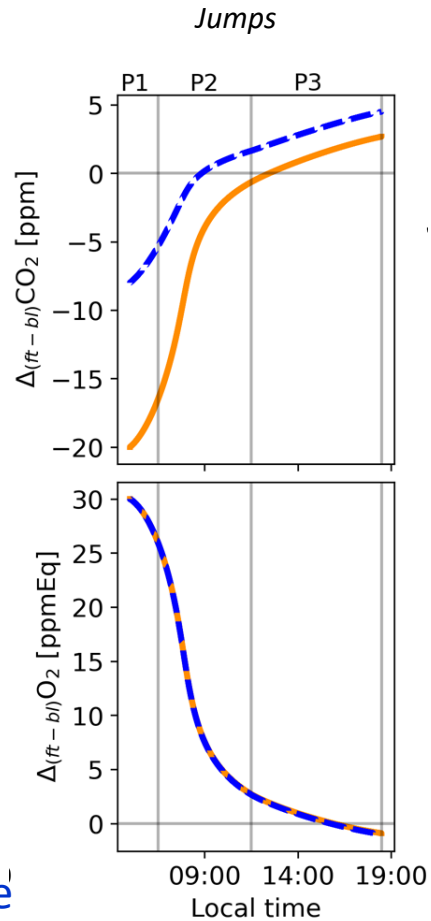
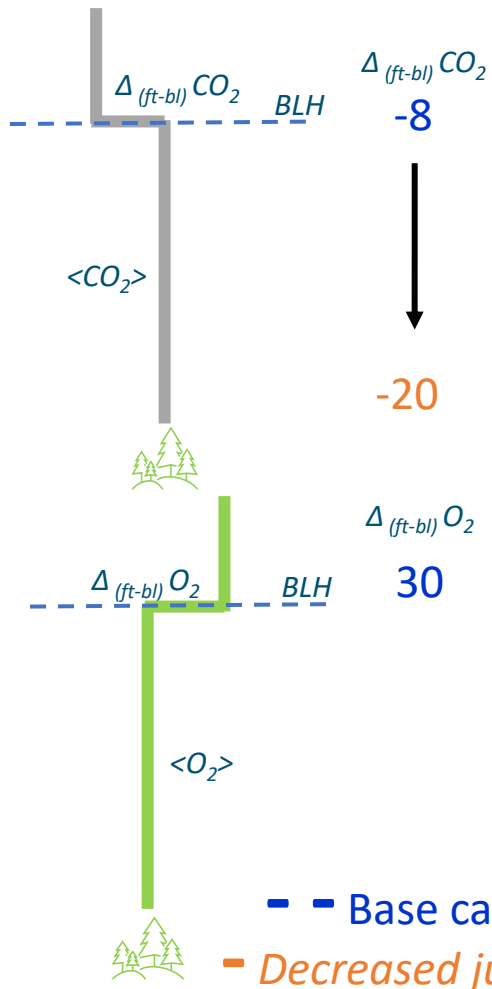
$$ER_{atmos} = ER_{forest} * \left( \frac{1 + \beta_{O_2}}{1 + \beta_{CO_2}} \right)$$

$$\beta = \frac{F_e}{F_s} = \frac{w_e * \Delta_{(ft-bl)}}{F_s}$$



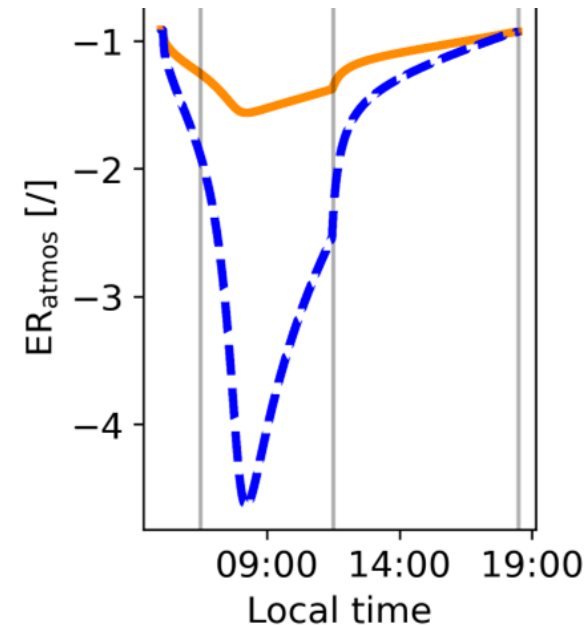
*P2: still high  $ER_{atmos}$  values*  
*P3:  $ER_{atmos}$  closer to  $ER_{forest}$*

# Other option: decrease jump ratio ( $\Delta O_2 : \Delta CO_2$ )



$$ER_{atmos} = ER_{forest} * \left( \frac{1 + \beta_{O_2}}{1 + \beta_{CO_2}} \right)$$

$$\beta = \frac{F_e}{F_s} = \frac{w_e * \Delta_{(ft-bl)}}{F_s}$$



P2: now lower  $ER_{atmos}$  values  
 P3:  $ER_{atmos}$  a bit closer to  $ER_{forest}$



# Take home message

*BL processes are of importance to understand  $O_2$  and  $CO_2$  mole fractions and their ratios*

## $ER_{forest}$ : based on 2/more heights

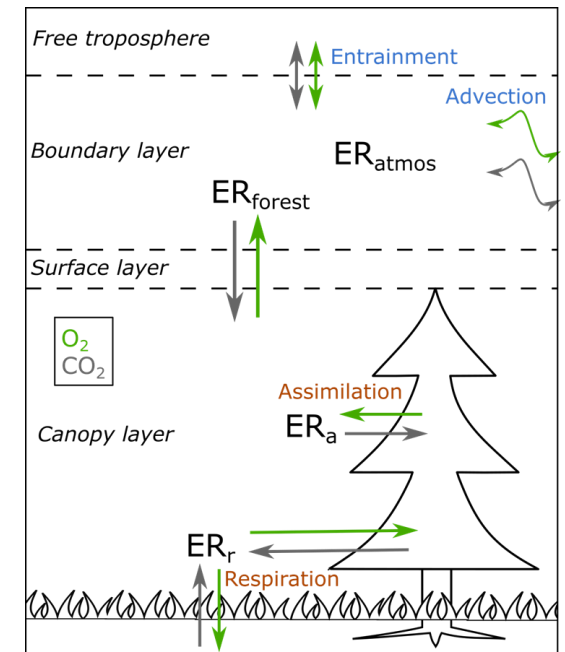
- This signal should be used to indicate surface
- $O_2$  fluxes can be calculated with the flux-gradient method
- New ER signals for boreal forest

Faassen et al. (2023)

## $ER_{atmos}$ : based on 1 height

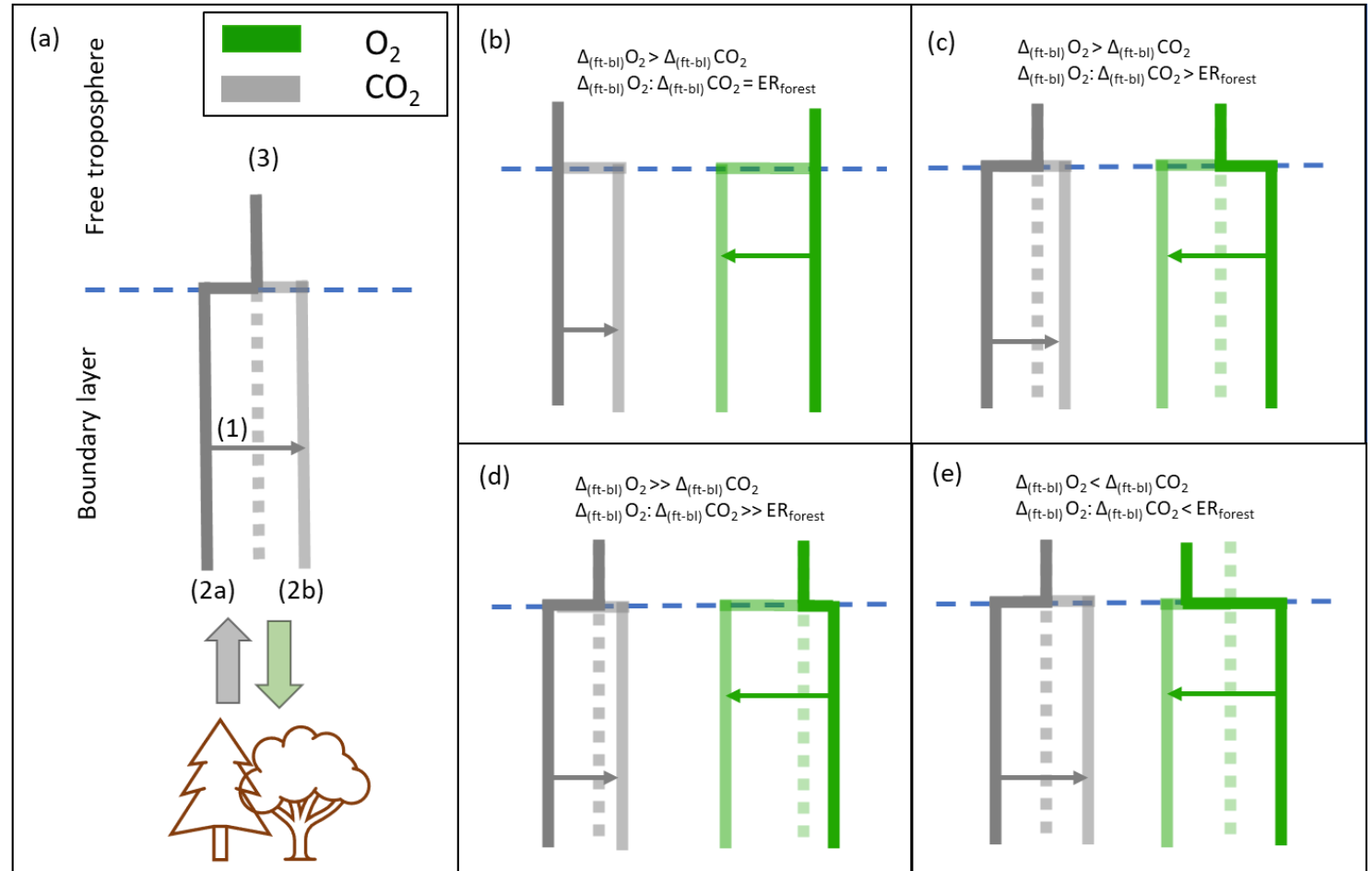
- Give more insights in non-local processes
- Handled with care
- Can get 'strange' values
- Budgets should be look at separately

Faassen et al. (in prep)





# Jump options



(1) Surface flux size during night (2a) mixed layer mole fraction just before sunset (2b) mixed layer mole fraction just before sunrise (3) Mole fraction in the free troposphere

# Flux calculations

